



# Results of Sediment Coring at the Town of Peace River, Northwestern Alberta (NTS 84C)

**Results of Sediment Coring at  
the Town of Peace River,  
Northwestern Alberta  
(NTS 84C)**

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## **Abstract**

The Peace River Lowlands of Alberta and British Columbia is one of the most historically active mass-movement areas in western Canada. Alberta Geological Survey (AGS), in collaboration with other stakeholders, has initiated a multiyear study, centred on the town of Peace River, to develop a better understanding of the type and extent of landsliding within the municipality. To help establish the geological setting for these landslides, AGS conducted a drilling program using rotosonic coring, wet-rotary drilling and downhole geophysics.

This report details the drill program and contains borehole logs with descriptions of the encountered lithology. The lithological descriptions are accompanied by geophysical logs and selected core photographs from two of the borehole locations. In addition, the report defines the methodology developed to classify discontinuities in the core samples.

# 1 Introduction

The Peace River Lowlands of Alberta and British Columbia are one of the most historically active mass-movement areas of Western Canada. Mass-movement events have affected, and continue to affect urban centres, such as the town of Peace River in northwestern Alberta, and their associated infrastructure. To better our understanding of these events, Alberta Geological Survey (AGS), in collaboration with various stakeholders, has initiated a multidisciplinary study of urban geology and landslide hazards (Froese, 2007; Morgan et al., 2008) of an area centred on the town of Peace River (Figure 1). This study aims to provide a better understanding of the type and extent of landslides within the Peace River valley. Data obtained from this investigation will be used to aid in municipal planning strategies that will reduce the risk of landslide hazards to population and infrastructure within the study area.

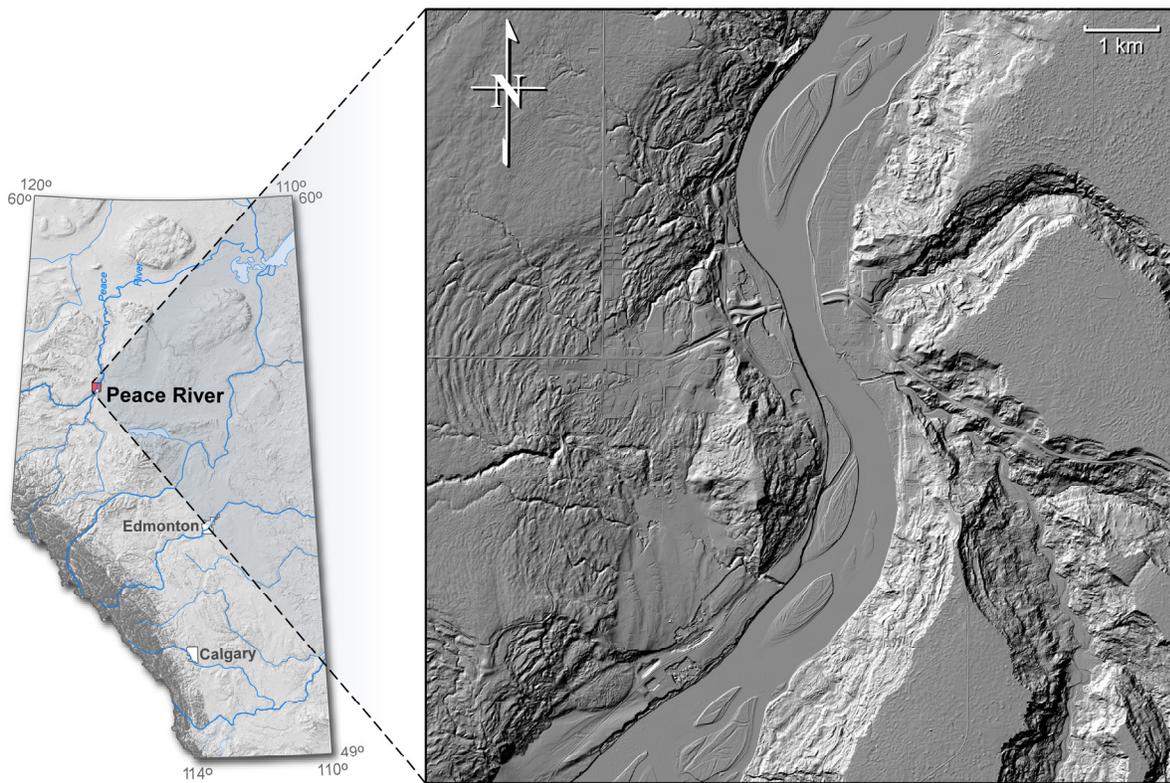


Figure 1. Location of study area.

This report provides a description of an AGS drilling program carried out as part of the landslide study in the town of Peace River in December 2008. The report briefly outlines the investigative methods used and provides the results of the drill program. We will incorporate the data from this program into the larger Peace River Urban Geology and Landslide Hazard Study.

## 2 Location and Physiography of the Study Area

The study area lies within the Peace River Lowlands physiographic zone of the Interior Plains (Bostock, 1981; Pettapiece, 1986; NTS 84C; Figure 1). Much of the surface morphology in the study area is the result of processes associated with the last glacial event (Late Wisconsinan) and Holocene erosion. The Peace River valley, the predominant morphological feature of the region, developed from incision of the Peace River through the Quaternary sediments and into the Cretaceous bedrock.

The river valley is an important transportation corridor and includes highways and a railway line. The Peace River Lowlands separates the Buffalo Head Hills to the east and the Whitemud and Clear Hills to the west. The Peace River District is composed mainly of fertile farmland, characterized by flat topography and stone-free fields developed on sediments deposited within Glacial Lake Peace, which inundated most of the region below 610 m above mean sea level (m asl; Mathews, 1980; Leslie and Fenton, 2001; Paulen, 2004; Paulen et al., 2004). Elevation within the study area varies from 320 m asl at the floodplain in the town of Peace River to a maximum of 625 m asl in the northwestern part of the study area.

### **3 Purpose**

To better our understanding of the type and extent of landslides within the Peace River Lowlands, we constructed a three-dimensional model using approximately 1400 available drillholes in the study area. The data consisted of logs from shallow geotechnical test holes and water wells, and downhole geophysical profiles obtained from oil and gas wells. As outlined in Morgan et al. (2008), these data, along with field mapping, were used to generate a schematic stratigraphy of the subsurface Quaternary sediments and the topography of the bedrock surface. To provide a greater level of detail on the lithology and sedimentology of the thick Quaternary sediments, a series of sediment cores, complemented by downhole geophysical surveys, was completed in and around the town of Peace River. Data obtained from this investigation will

- improve upon existing stratigraphic modelling of Quaternary and Holocene sediments of the Peace River valley;
- aid in determining the potential mechanism(s) for landslide activity in the Peace River valley;
- aid in determining the type of landslide events that occurred in the Peace River valley;
- aid in determining the deglacial history of the Peace River valley;
- provide high-quality subsurface data that will be used as calibration points during the construction of a geological model of the Peace River valley;
- provide representative sections of the subsurface stratigraphy to aid the geotechnical consulting community in identifying both intact and displaced units within the subsurface;
- provide a spatial distribution of subsurface lithologies in the Peace River area; and
- supply all levels of government, the public and the scientific community with an improved understanding of landslide activity in the Peace River valley that will aid in land-use and infrastructure planning strategies.

### **4 Methodology**

This section provides a brief description of drilling and geophysical methods used in this study. For a more detailed assessment of overburden-drilling methodology and downhole geophysical techniques, see Allen (1993) and Schlumberger (1991), respectively.

#### **4.1 Drilling Methods**

The Peace River drill program used two drill rigs, each of which employed a different drilling technique. A truck-mounted Ingersoll Rand T2W rig drilled mud-rotary boreholes through both soil and bedrock, and a truck-mounted Ingersoll Rand TH-60 sonic rig collected continuous core samples through the upper 114 m of the soil column. Brief descriptions of the two drill methods follow.

### **4.1.1 Mud-Rotary**

This drilling method uses a tricone drill bit on the bottom of a string of drill rods that rotate within the borehole. Drilling fluid is pumped into the borehole through the string of drill rods, thereby flushing the drill cuttings (sediment and bedrock samples) to the surface via the space between the drill rods and the borehole wall (annulus). Examination of the returned cuttings is the primary source of lithology determination, although changes in the rate of drilling progress (more commonly called rate of penetration or ROP) or pump pressure can indicate a change in material consistency or lithology, and are commonly helpful as indicators of material boundaries when interpreted by an experienced driller.

### **4.1.2 Sonic**

The sonic method (also termed roto-sonic) uses vibration as a means for tool advancement. A hydraulically actuated drilling head is employed to manipulate an inner casing (core barrel) and an outer casing, provide a vibratory source and allow for rotation of the casing. The drilling head applies a sinusoidal vibration to the inner casing, which has a cutting head attached at its end, while at the same time pushing the casing into the ground, typically in 3.05 or 6.10 m (10 or 20 foot) intervals. The outer casing then advances until it reaches the depth of the inner casing, to maintain the borehole. Drilling proceeds in this way, with additional casing segments added as hole depth increases. Once the desired depth is reached, the inner casing is extracted from the hole and the continuous core sample is removed, or vibrated, from the core barrel. A geologist or technician typically examines the core sample on site before it is packed and shipped to a laboratory or discarded, depending on the application.

## **4.2 Drilling Locations**

The AGS Peace River drill program was completed between November 25 and December 16, 2008. Six boreholes were drilled at four sites using both conventional mud-rotary and sonic drilling methods. Figure 2 shows the locations of the boreholes.

At both of the sonic borehole sites, a mud-rotary borehole was drilled approximately 10 m away from the sonic hole. The sonic rig was used to sample the Quaternary sediments overlying the Cretaceous bedrock to the maximum feasible depth based on drilling conditions, whereas the corresponding mud-rotary hole was drilled through the entire sequence of Quaternary sediments and into the Cretaceous bedrock below. These mud-rotary holes were used to ascertain the lithology of the bedrock beneath the adjacent sonic borehole and to facilitate the collection of a suite of downhole geophysical logs, which requires a fluid-filled hole.

Table 1 lists the type, depth, location and order in the sequence of each hole, and indicates the number of the relevant borehole log in Appendix 1.

## **4.3 Sample Description and Logging**

During drilling operations, a geological technologist, sedimentologist or geological engineer was present to provide direction to the rig crew and to receive, examine and describe the drill samples. During mud-rotary operations, soil samples (drill cuttings) were collected at 1.5 m (5 ft) intervals. For sonic operations, the sample interval was effectively continuous, with 5.4 and 2.5 m of lost core reported over 108.2 and 114.3 m of coring in holes PR08-3 and PR08-5, respectively. For both sonic holes, 3.05 m (10 feet) core runs were made to a depth of 15.24 m (50 feet) below ground surface, after which core runs of 6.10 m (20 feet) were made. Due to the expansive nature of the overconsolidated silts, clays and tills that were being sampled, the driller used a drilling mud to lubricate the core barrel and casing during drilling to minimize the chance of getting them stuck in the hole.

The mud-rotary cuttings and the sonic core samples were described on site based on all or some of the following properties: texture, colour, clast lithology, percentage of pebbles and granules, relative

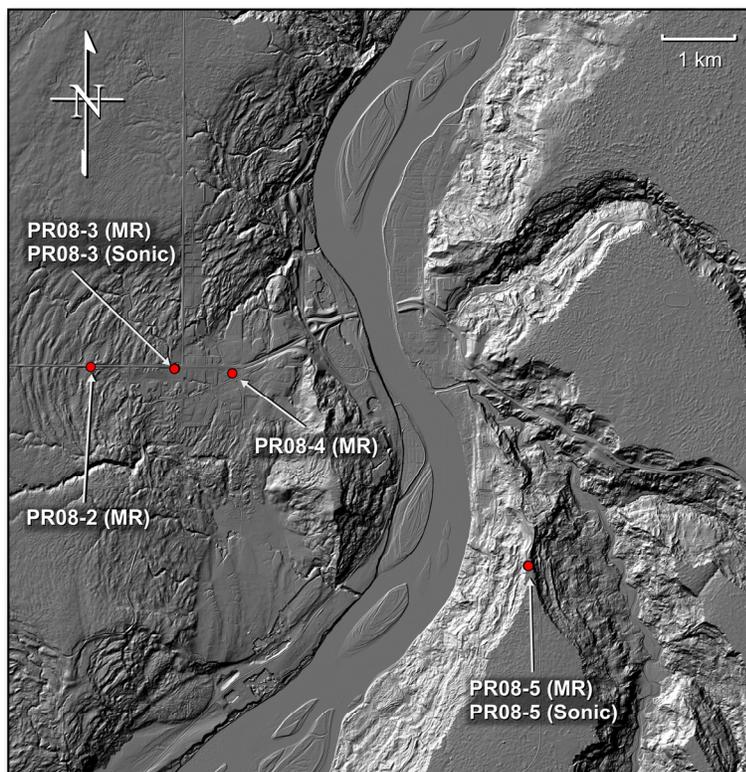


Figure 2. LiDAR image showing locations of mud-rotary (MR) and sonic boreholes.

moisture, and structure. Colour was based on a standardized soil colour chart (Munsell, 1990) and physical soil properties were based on standard methods outlined by Folk (1980). Other descriptive comments included sediment structure (e.g., debris flow, load structures, layering, etc.) and the presence of other relevant features, such as fractures or slickensides. Appendix 2 provides photos of selected intervals and features within the sonic core. After examination and description of the samples in the field, the sonic core was placed in PVC split casing and the mud-rotary cuttings were sealed in plastic bags. All the samples were shipped to the AGS Mineral Core Research Facility (MCRF) in Edmonton for further detailed analysis.

Table 1. Borehole types and locations.

Borehole Name & Type	Total Depth (m)	UTM Location (Zone 11N, NAD83)	Order Drilled	Borehole Log ID (Appendix 1)
PR08-2 Mud-Rotary	172.2	477235E, 6232069N	3 <sup>rd</sup>	PR08-2
PR08-3 Mud-Rotary*	125	478380E, 6232042N	5 <sup>th</sup>	PR08-3 COMP
PR08-3 Sonic	108.2	478388E, 6232046N	2 <sup>nd</sup>	
PR08-4 Mud-Rotary	30.5	479169E, 6231983N	6 <sup>th</sup>	PR08-4
PR08-5 Mud-Rotary*	177	483207E, 6229415N	1 <sup>st</sup>	PR08-5 COMP
PR08-5 Sonic	114.3	483199E, 6229419N	4 <sup>th</sup>	

\* Denotes holes where downhole geophysical logs were run.

We finalized the logs (Appendix 1) for the mud-rotary holes after ‘truthing’ the field logs by examining and confirming the lithology of the preserved samples at the MCRF. For those holes on which we conducted downhole geophysics, we used the resulting geophysical logs to help pick lithological boundaries and provide information between samples.

Two of the logs in Appendix 1 are in a composite log format, which displays the entire depth of the sonic corehole with associated lithological description and then continues using the lithology determined from the mud-rotary borehole. The geophysical logs are also displayed against the lithology information from the sonic and mud-rotary holes.

We re-examined the sonic core in detail at the MCRF to confirm the field logging and create expanded descriptions of sedimentological and geotechnical features. Discontinuities or shearing within the core samples received special attention during logging, particularly as zones of weakness such as these within Quaternary sediments can be indicative of movement zones associated with landsliding in the region. We described the discontinuities in the core using a classification system described in Appendix 3, which was designed for logging of geotechnical discontinuities in mudstones and shales. This classification allows for a systematic description of each discontinuity based on parameters such as spacing, continuity, core breakage, geometry, roughness, sheen and orientation.

#### **4.4 Geophysics**

Wireline geophysical measurements were collected in two mud-rotary holes, PR08-3 and PR08-5, which were drilled adjacent to the two sonic coreholes. The suite of wireline services run included standard gamma-ray, spontaneous-potential (SP), single-axis caliper, focused-electric (resistivity) and density measurements. The plots of these measurements are included in the composite logs in Appendix 1. This section provides a brief description of each measurement (Schlumberger, 1991).

The standard gamma-ray reading is a recording of the natural radioactivity of the formation. The gamma-ray tool uses a detector to measure the gamma rays that are spontaneously emitted from radioactive elements, nearly all of them coming from the potassium, uranium and thorium series. In sedimentary formations, the log reflects the clay or shale content because radioactive elements tend to concentrate in clays and shales. Clean formations (e.g., sand, sandstone, carbonates) typically have low levels of radioactivity unless they contain naturally occurring radioactive mineral grains.

The spontaneous-potential (SP) log is a measurement system that records the electric potential (voltage) produced by the interaction of formation connate water, conductive drill fluid and certain ion-selective materials. The SP plot is a recording versus depth of the difference between the electrical potential of a moving electrode and that of a surface electrode. Opposite shales or clays, the plot usually defines a generally straight line. Opposite permeable formations, it shows deviations from this straight line. The SP plot is useful for differentiating the porous and permeable formations (sands, sandstones, limestones) from the nonpermeable clays and shales. The SP measurement also has other uses more applicable to oil field applications, such as the determination of relative salinity and resistivity of formation water.

Focused-electric measurement provides a value for the resistivity, or specific resistance, of the formation. Simply put, the resistivity device emits a current of constant intensity between two current electrodes and measures the resultant potential difference between two measurement electrodes on the tool. Resistivity is determined based on these measured voltages. The wireline tool configuration used in the Peace River project provided a relatively shallow reading of resistivity (i.e., close to the borehole wall).

Density tools provide a measurement of the bulk density of the formation. A radioactive source emits medium-energy gamma rays into the formation. As the gamma rays collide with the electrons in the formation, they lose energy. These scattered gamma rays reach detectors located at fixed distances from the source and are counted as an indication of formation density. Density tools typically have a caliper

arm integrated into the tool. In addition to providing a measurement of borehole diameter, the caliper forces the density skid against the borehole wall for optimal measurement when opened.

The composite logs for boreholes PR08-3 and PR08-5 (Appendix 1) include the geophysical plots beside the lithology descriptions.

#### **4.5 Drilling Conditions**

Because the Peace River drill program was during the winter, it had to contend with operational delays typical of cold-weather operations. The drilling program achieved the desired results with a few exceptions that were the result of normally encountered drilling issues.

Two of the mud-rotary boreholes encountered borehole stability problems. Borehole PR08-2 advanced successfully to a depth of 172 m below surface, 6 m into the Cretaceous shale bedrock underlying the Quaternary sediments. As the crew was tripping the drillstem out of the hole, the borehole began to collapse about 28 m below ground surface in the sand overlying the diamicton (till). Attempts were made to condition the borehole in preparation for the geophysical wireline runs, but hole integrity worsened and the borehole had to be abandoned and the geophysical logging cancelled.

Borehole PR08-4 was drilled at the end of the program, with limited operational budget remaining. Poor drilling conditions were encountered at 30 m below ground surface in sand and gravel deposits within the diamicton (till, based on drilling response). Attempts to stabilize the borehole were unsuccessful and the decision was made to abandon the hole, as spending had reached the program's budgetary limit.

Sonic hole PR08-3 reached a maximum core depth of 108 m. The drill rig lost power due to engine problems and required 1.5 days for repairs. During this downtime, the outer casing became stuck in the hole and the borehole had to be abandoned after retrieval of all but some of the lower outer casing.

### **5 Summary and Conclusions**

The program drilled and cored a series of five test holes at the town of Peace River. We chose the locations to provide detailed stratigraphic and geotechnical information in both intact ground and landslide terrain throughout the study area, and to determine depth to bedrock as verification for the previously prepared 3-D geological model. The two sonic boreholes, PR08-3 and PR08-5, provided continuous core for the upper 108 and 114 m, respectively, of the stratigraphy. These complemented mud-rotary drilling and downhole geophysical testing into the top of bedrock (more than 170 m below the surface). We expect these test holes will not only aid in interpretation of the Quaternary stratigraphy and bedrock surface orientation, but will also be a significant reference for professionals planning geotechnical investigations and designing mitigatory measures for landslides in the Peace River region.

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## **Appendix 1 — Borehole Logs**

PR08-2 (page 9)

PR08-3 (page 17)

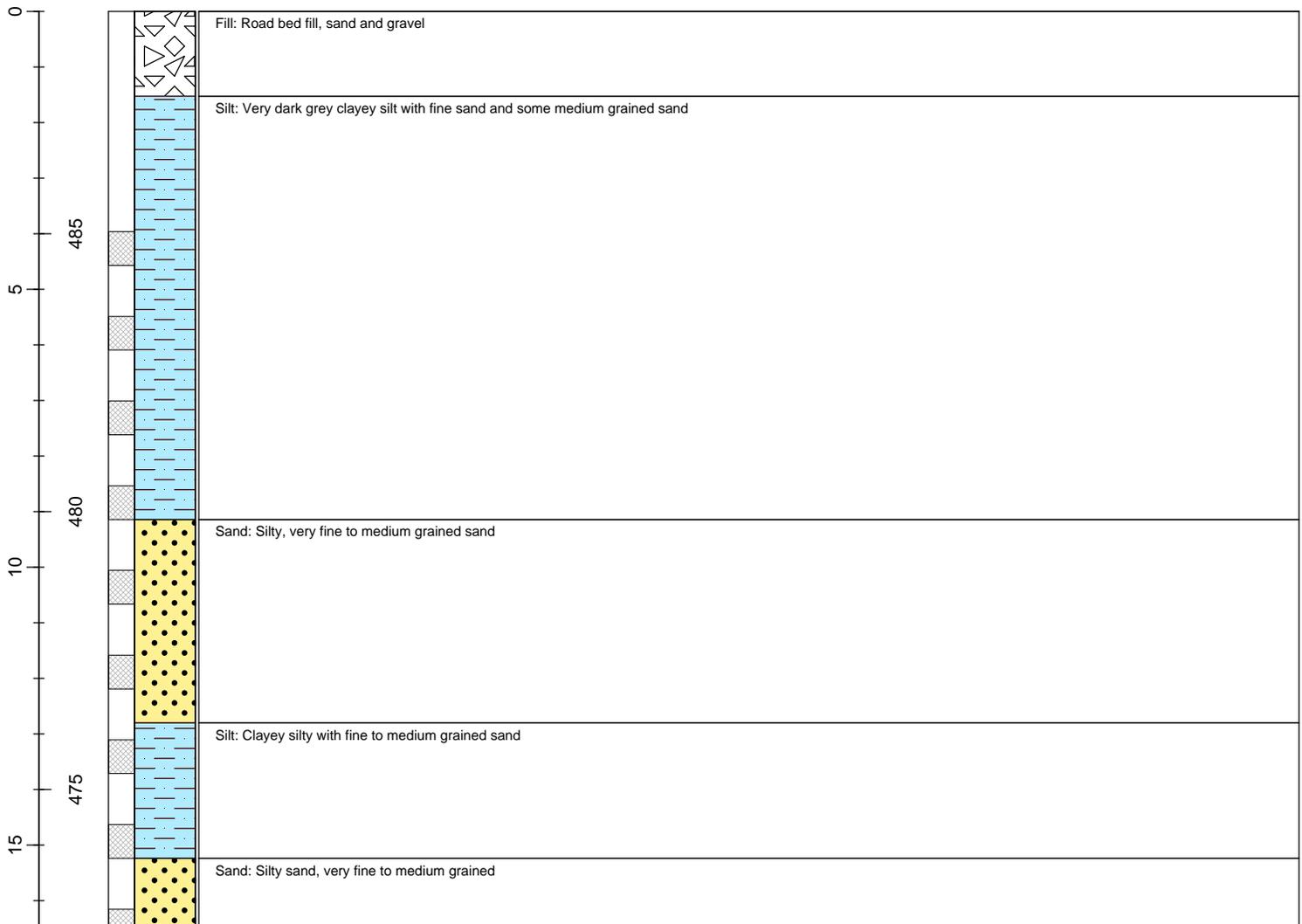
PR08-4 (page 22)

PR08-5 (page 24)

Log ID: **PR08-2**

Total Depth: **172.2 m**  
 Location: **113-26-083-22W5M**  
 Northing (UTM Nad83): **6232069**  
 Easting (UTM Nad83): **477235**  
 Hole Diameter: **155 mm**  
 Elevation (Ground Surface): **489 m asl**  
 Drilling Date: **12/10/08**  
 Drilled By: **Garrity and Baker (Wet Rotary)**  
 Lithology Logged By: **J. Warren/ S. Slattery**  
 Geophysical Log Operator: **N/A**

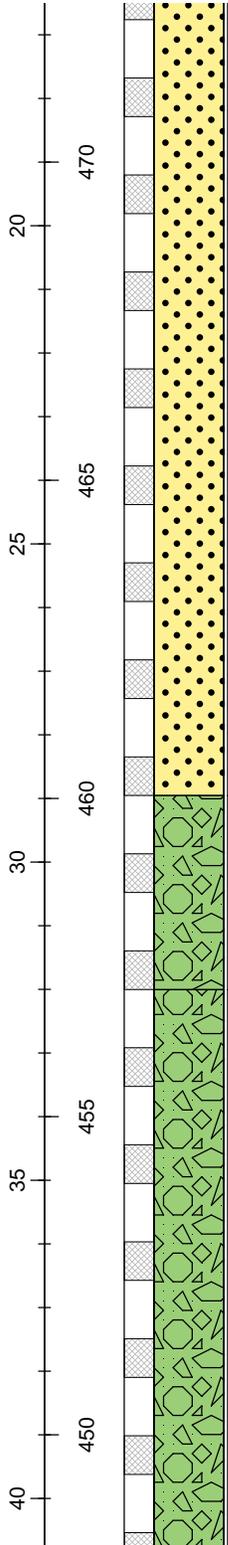
Depth (Meters)	Elevation (m asl)	Sample Interval	Lithology
----------------	-------------------	-----------------	-----------



Depth (Meters)  
Elevation (m asl)

Sample Interval

Lithology



Diamicton: Clayey silt till with very fine to medium grained sand, some fine gravel sizes

Diamicton: ...rocks in hole causing drilling/stability problems, driller cleaned hole.

Depth (Meters)  
Elevation (m asl)

Sample Interval

Lithology

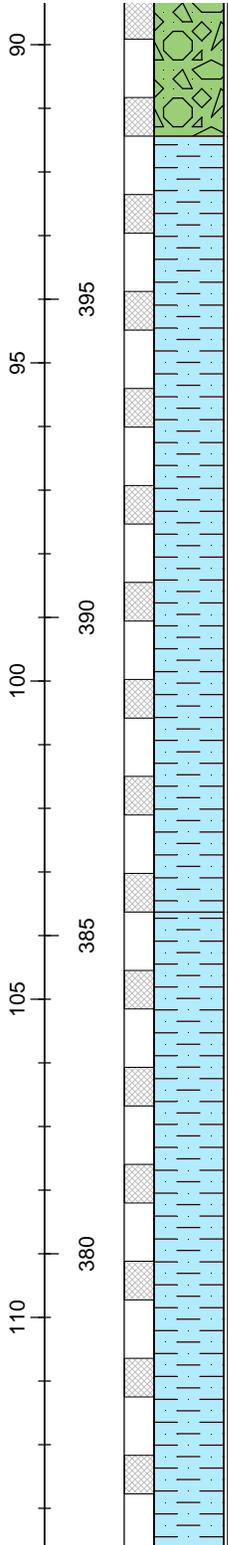




Depth (Meters)  
Elevation (m asl)

Sample Interval

Lithology



Silt: Drilling response indicative of a lithology change to silts and clays. Very dark grey clayey silty with minor fine grained sand.

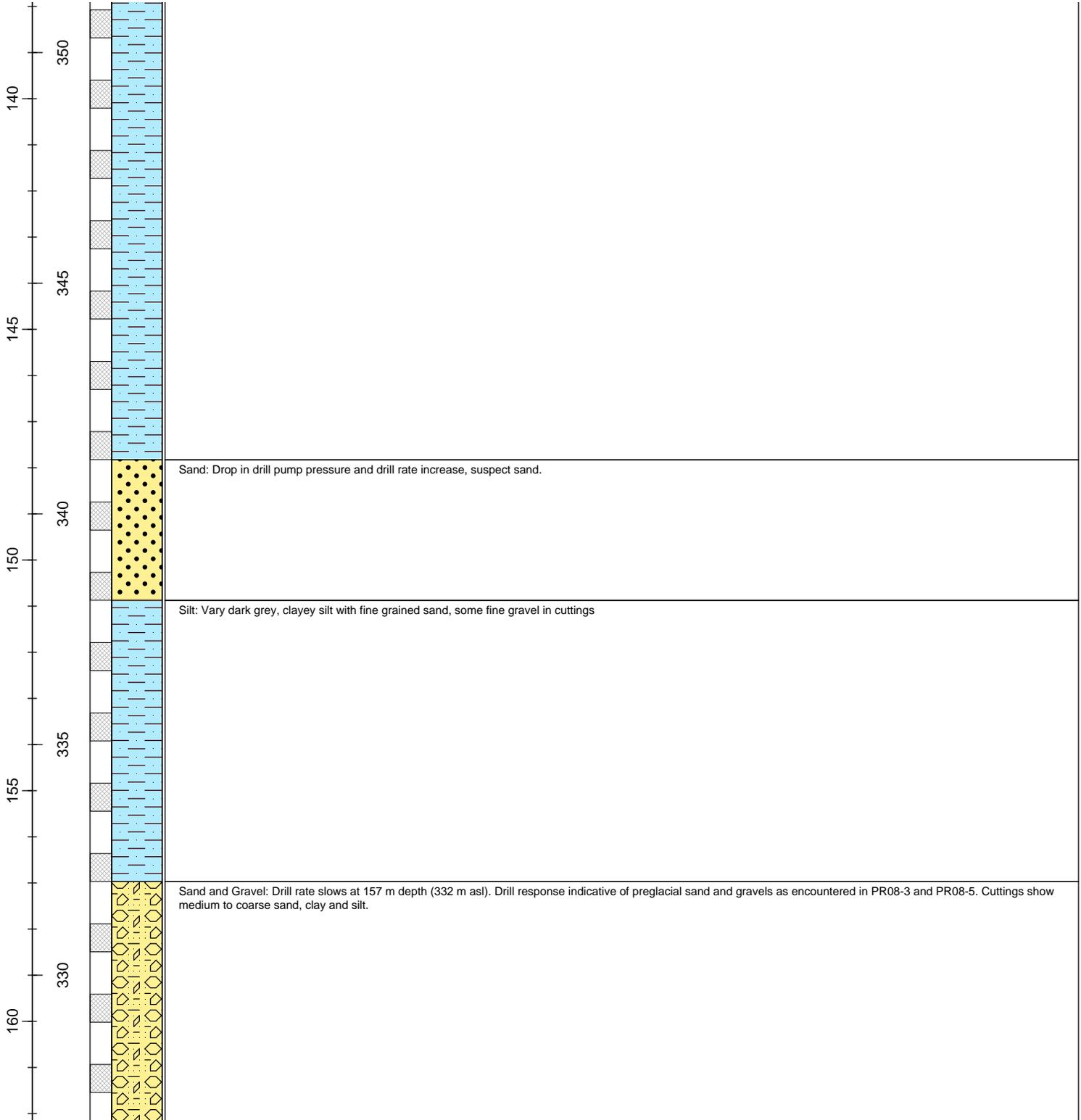
Silt: Drilling slower at 103.6 m (385.4 m asl). Very dark grey clayey silt with fine grained sand, some fine gravel in cuttings.



Depth (Meters)  
Elevation (m asl)

Sample Interval

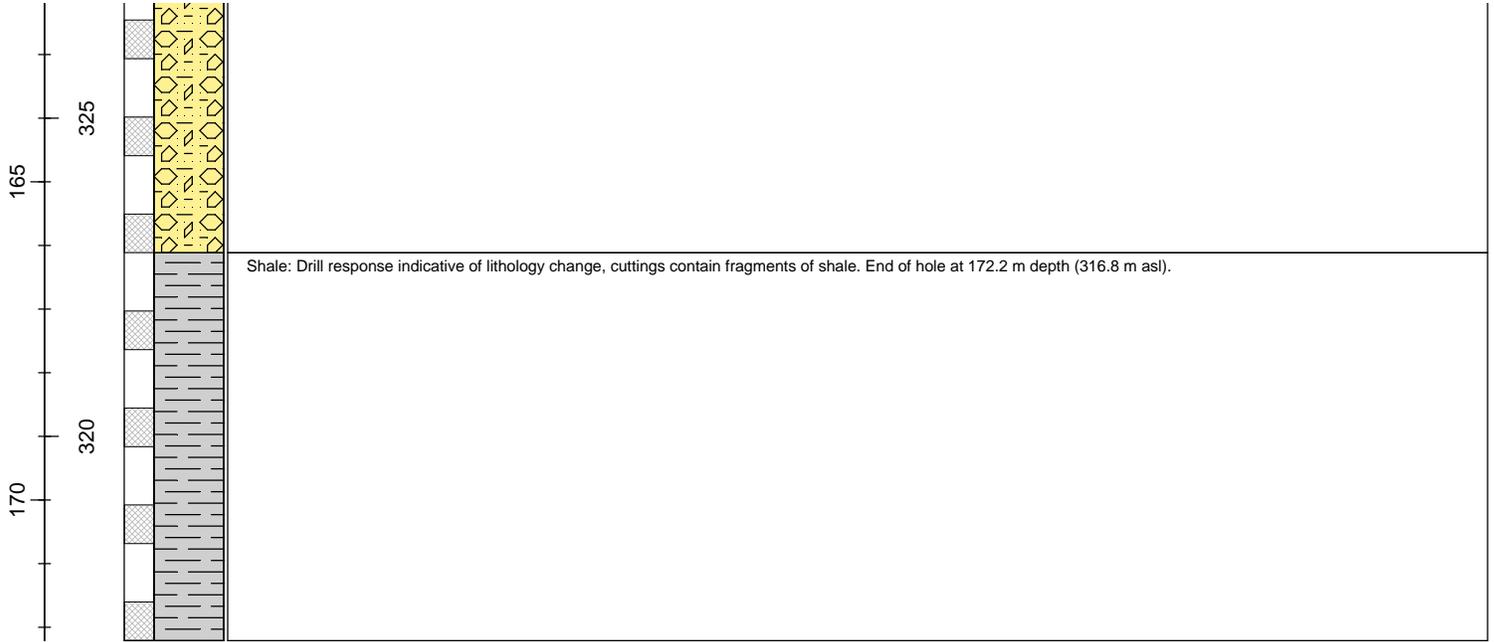
Lithology



Depth (Meters)  
Elevation (m asl)

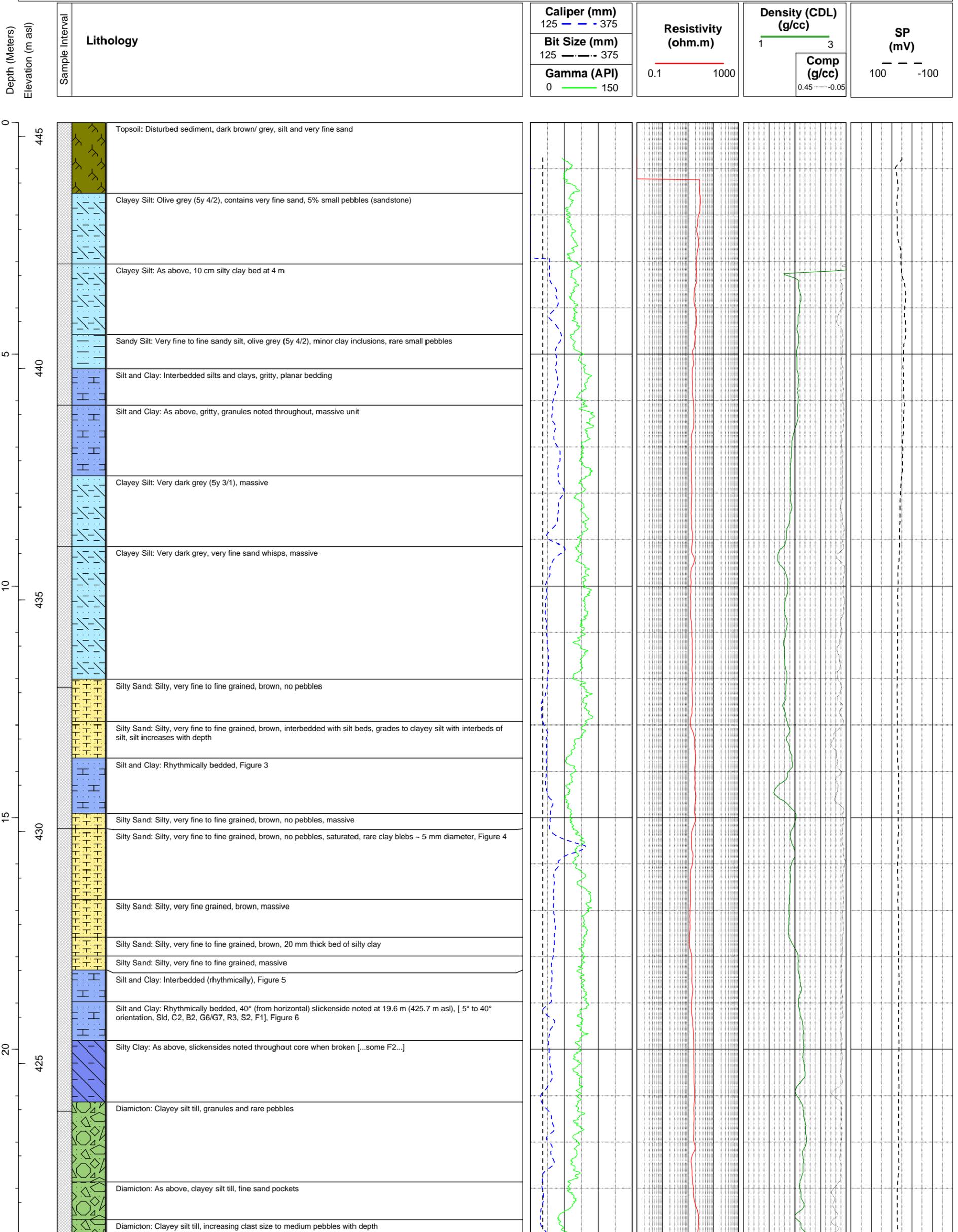
Sample Interval

**Lithology**



Log ID: **PR08-3 Composite Log**

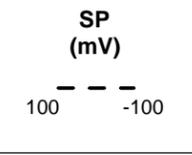
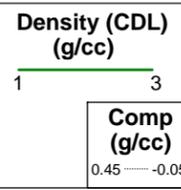
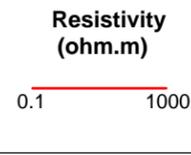
Total Depth: **124.97 m**  
 Location: **16-26-083-22W5M**  
 Northing UTM Nad83: **6232046**  
 Easting UTM Nad83: **478388**  
 Hole Diameter: **155 mm**  
 Elevation (Ground Surface): **445.3 m asl**  
 Drilling Date: **12/12/08**  
 Drilled By: **Boart Longyear (Sonic)/ Garrity Baker (Wet Rotary)**  
 Lithology Logged By: **S. Slattery/ J. Morgan/ J. Warren**  
 Geophysical Log Operator: **Century Wireline Services/ D. Sinkwich**



Depth (Meters)  
Elevation (m asl)

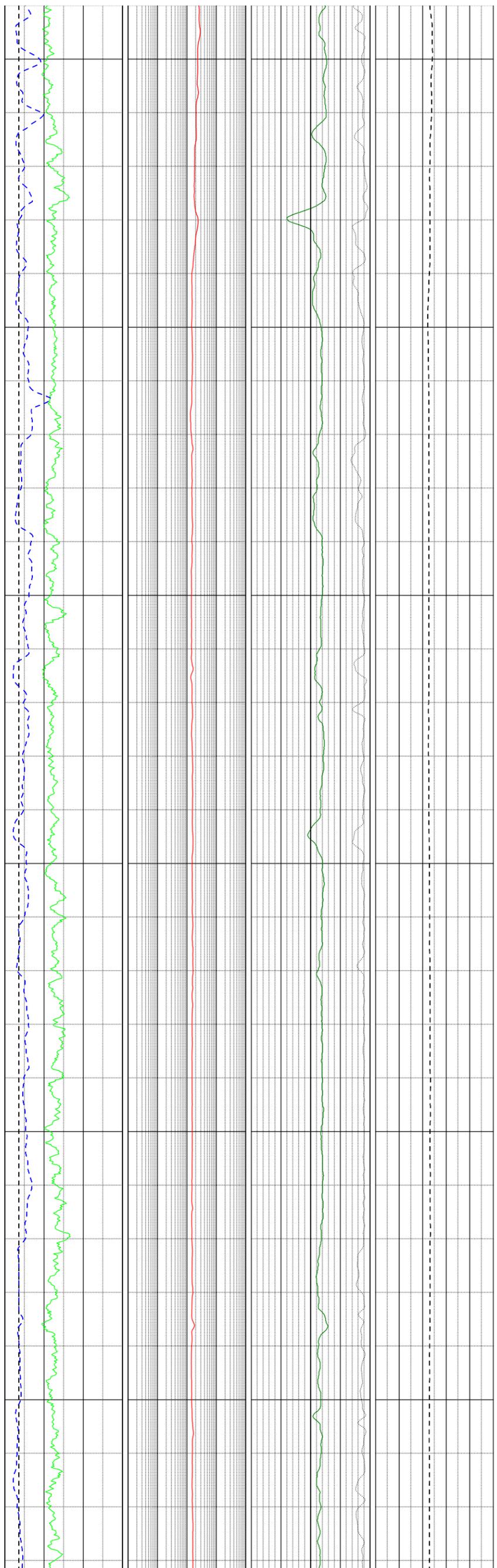
Sample Interval	Lithology
-----------------	-----------

<b>Caliper (mm)</b> 125 — 375
<b>Bit Size (mm)</b> 125 — 375
<b>Gamma (API)</b> 0 — 150



25  
30  
35  
40  
45  
50

420	Diamicton: Clayey silt till, clast rich, sandy lenses at base of core, Figure 7
	Diamicton: Clayey silt till, clast rich, granites, sandstones, Precambrian lithologies
	Sand: Silty, fine to medium grained, no pebbles, Figure 8
	Sand: Fine to medium grained, 10 cm of pebbles on top, inverse graded
	Sand: Fine to medium grained, pebbly
	Sand: Fine to medium grained, interbedded with silt
	Diamicton: Clayey silt till, clasts moderate (small pebbles)
	No recovery: 2.36 m lost core, boulder push, possibly sand
415	
	Sand: Fine to medium grained
	Diamicton: Clayey silt till, clast content ~ 5%
	Sand: Fine to medium grained, rare medium grade pebbles
	Sand: Silty, fine grained, pebbly till
	Diamicton: Clayey silt till, clast content ~ 5%, no HCl reaction
	Diamicton: As above, sandstone boulder at 33.7 m (411.6 m asl), green coloured fine grained sand at 34.5 m
410	
	Diamicton: Clayey silt till, very dark grey, clast content ~ 5%, no HCl reaction
	Diamicton: Clayey silt till (as above), slightly sandy, clasts common, coal fragments
405	
	No recovery: Core lost
	Sand: Silty sand
	Diamicton: Clayey silt till, very fine to fine grained sandy, dark grey, consolidated (compact), clast content 5-10%
400	
	Diamicton: Very fine grained sandy, clayey silt till, dark grey, dense, clast content 5-10%
	Diamicton: As above, massive, contact
	Sand and Silt: Alternating beds of clayey, sandy silt and silty, very fine grained sand. Heavily contorted units; load structures/ flames evident throughout. Convolute beds around clasts. Contact between till and glaciolacustrine unit is gradational. Figures 9, 10, and 11, (possible hyperconcentrated suspension fallout - glaciolacustrine)
395	
	Diamicton: Contorted/ liquefied silty sand and sandy silt beds interbedded with dark grey, dense, gritty, silty till
	Diamicton: Silty till, contains very fine to fine grained sand, massive, dense. 30 mm beds of contorted sand at 51.7 m (393.6 m asl).
	Diamicton: Silt till with very fine grained sand, dark grey, dense, stone content ~ 5%



Depth (Meters)  
Elevation (m asl)

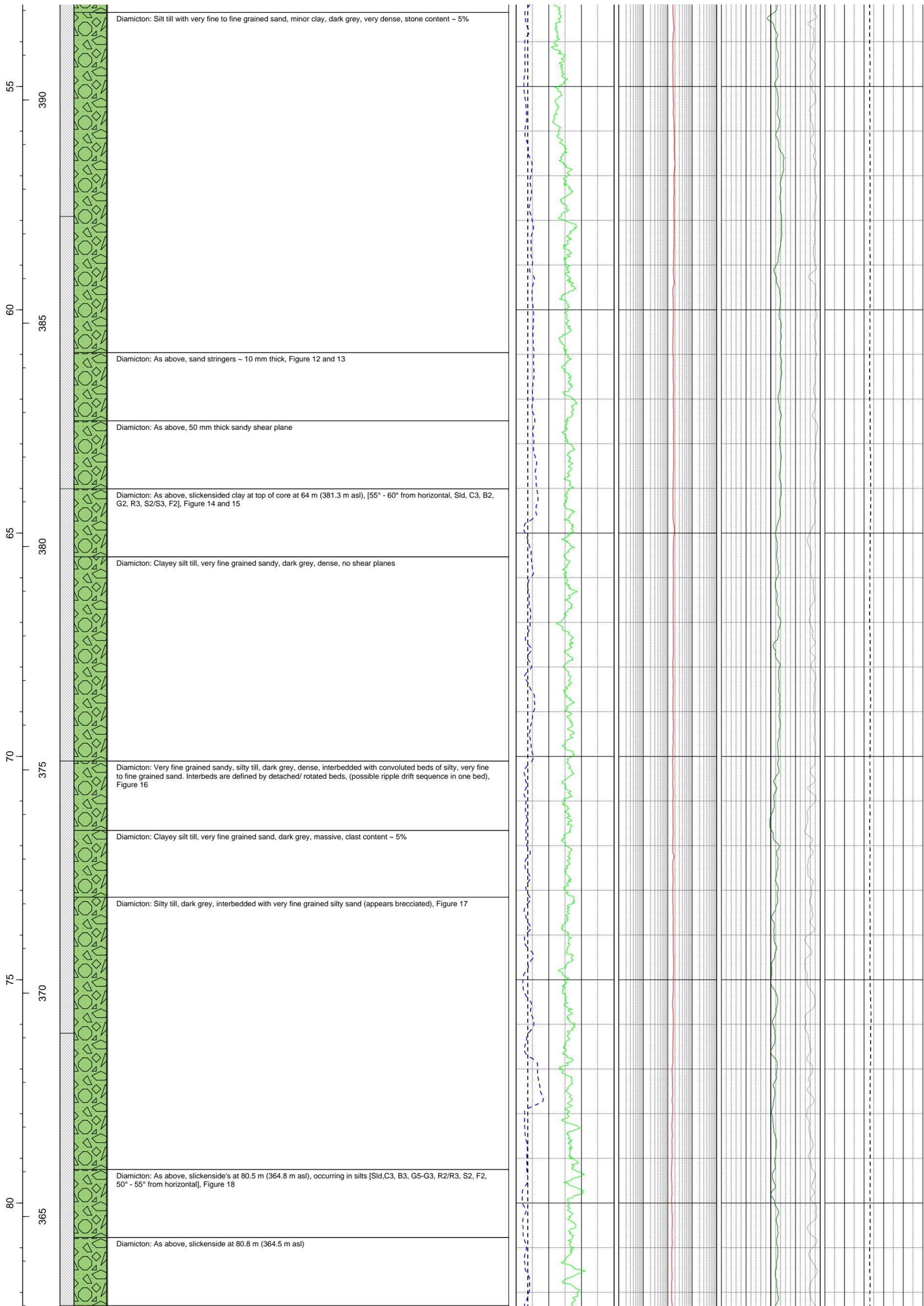
Sample Interval	Lithology
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<b>Caliper (mm)</b> 125 — 375
<b>Bit Size (mm)</b> 125 — 375
<b>Gamma (API)</b> 0 — 150

<b>Resistivity (ohm.m)</b>
0.1 — 1000

<b>Density (CDL) (g/cc)</b>
1 — 3
<b>Comp (g/cc)</b>
0.45 — -0.05

<b>SP (mV)</b>
100 — -100



Depth (Meters)  
Elevation (m asl)

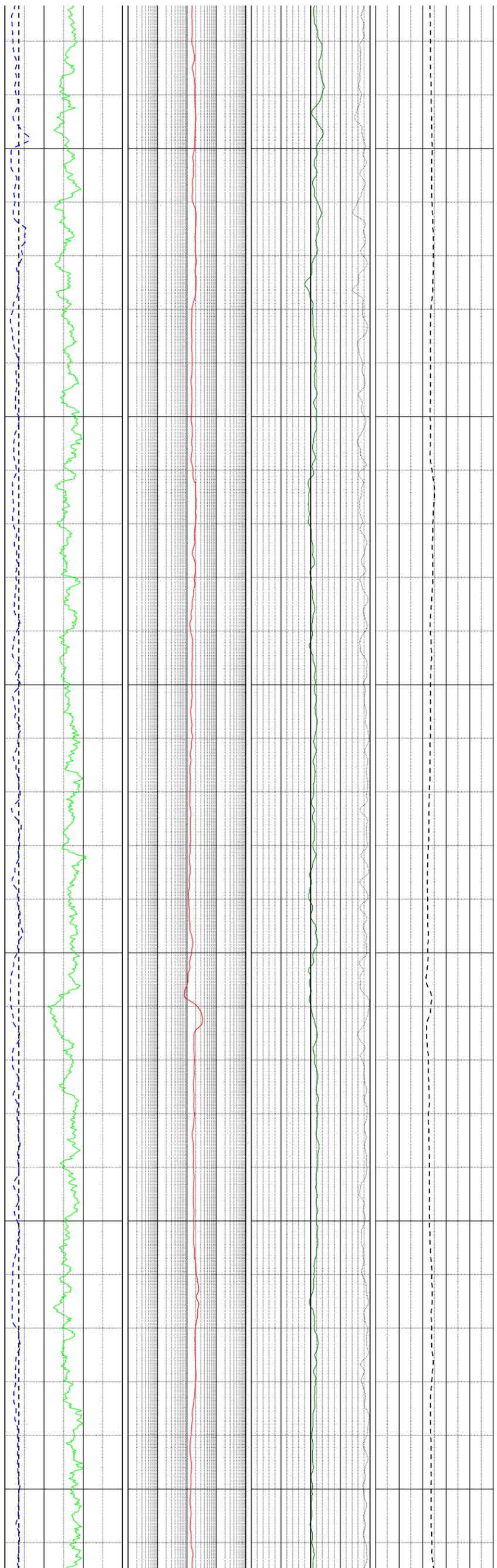
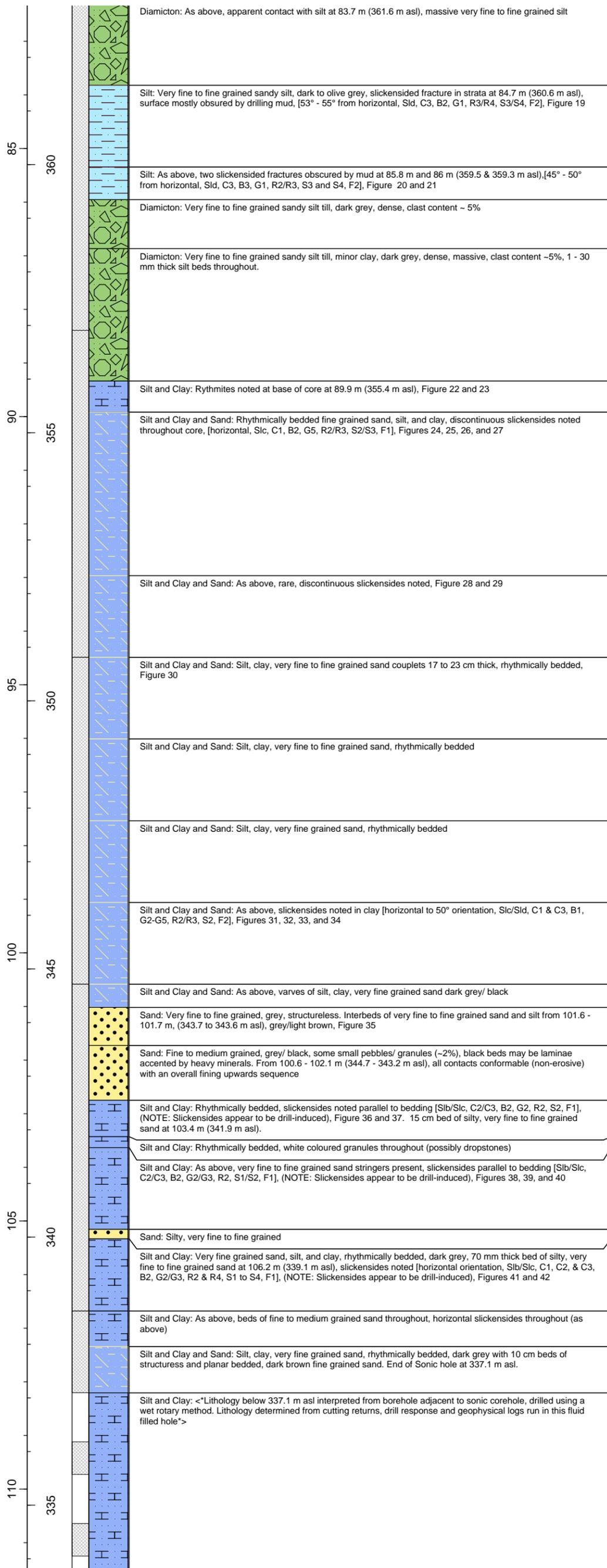
Sample Interval	Lithology
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<b>Caliper (mm)</b> 125 — 375
<b>Bit Size (mm)</b> 125 — 375
<b>Gamma (API)</b> 0 — 150

<b>Resistivity (ohm.m)</b> 0.1 — 1000
--

<b>Density (CDL) (g/cc)</b> 1 — 3
<b>Comp (g/cc)</b> 0.45 — -0.05

<b>SP (mV)</b> 100 — -100
------------------------------



Depth (Meters)  
Elevation (m asl)

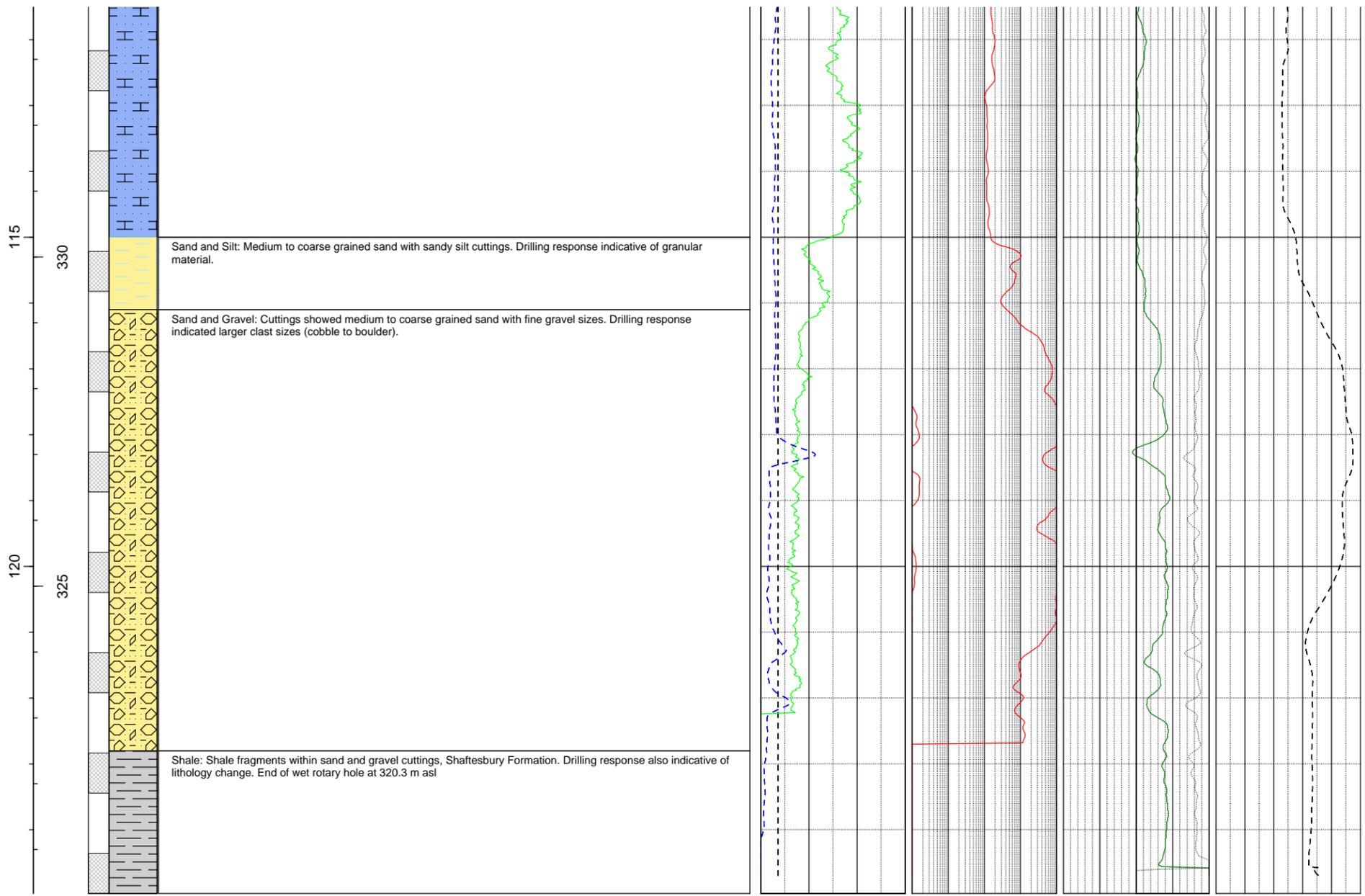
Sample Interval	Lithology
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<b>Caliper (mm)</b> 125 — 375
<b>Bit Size (mm)</b> 125 — 375
<b>Gamma (API)</b> 0 — 150

<b>Resistivity (ohm.m)</b>
0.1 — 1000

<b>Density (CDL) (g/cc)</b>
1 — 3
<b>Comp (g/cc)</b>
0.45 — -0.05

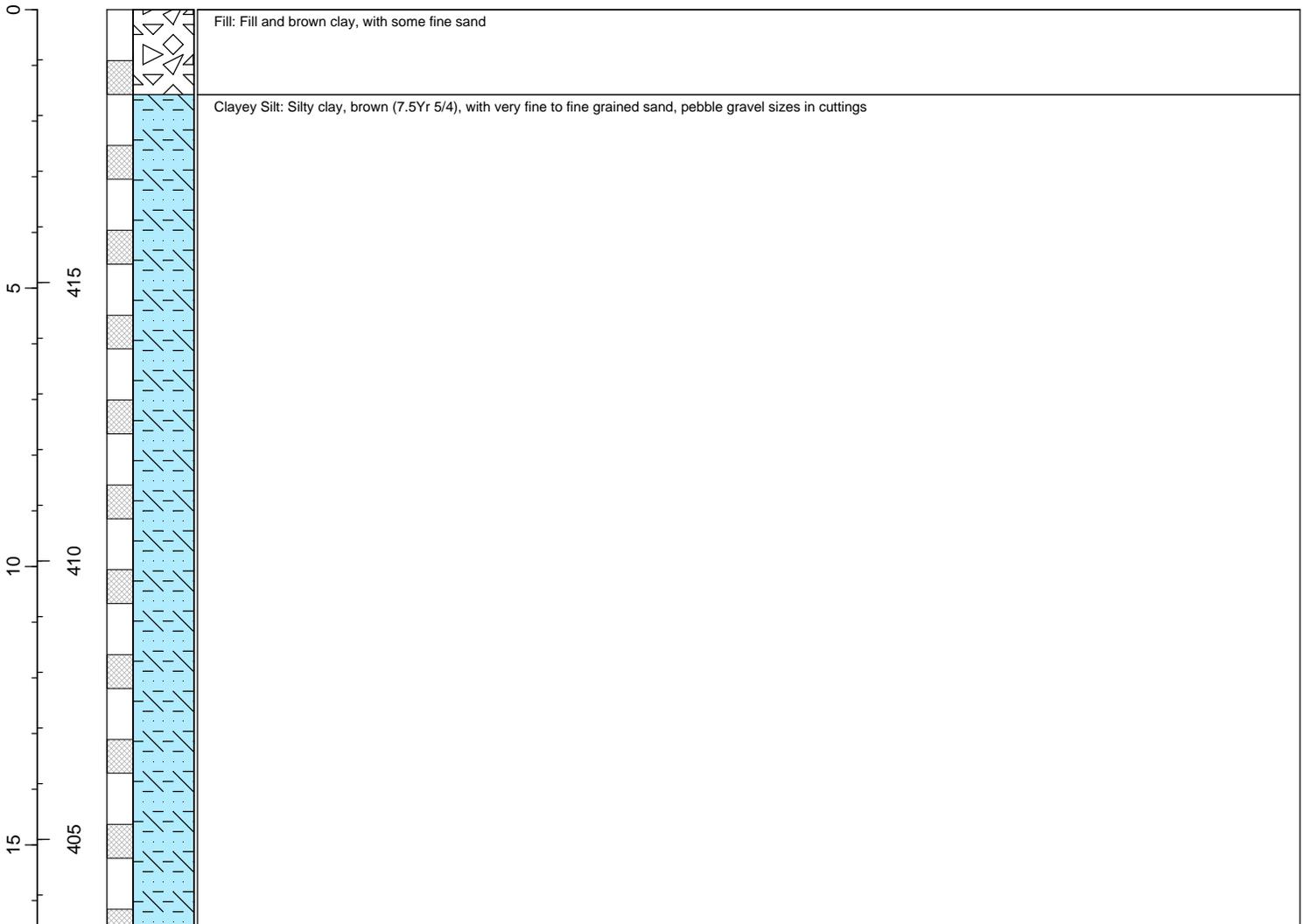
<b>SP (mV)</b>
100 — -100



Log ID: **PR08-4**

Total Depth: **30.5 m**  
 Location: **14-25-083-22W5M**  
 Northing (UTM Nad83): **6231983**  
 Easting (UTM Nad83): **479169**  
 Hole Diameter: **155 mm**  
 Elevation (Ground Surface): **419.9 m asl**  
 Drilling Date: **12/15/08**  
 Drilled By: **Garrity and Baker (Wet Rotary)**  
 Lithology Logged By: **J. Morgan**  
 Geophysical Log Operator: **N/A**

Depth (Meters)	Elevation (m asl)	Sample Interval	Lithology
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Depth (Meters)  
Elevation (m asl)

Sample Interval

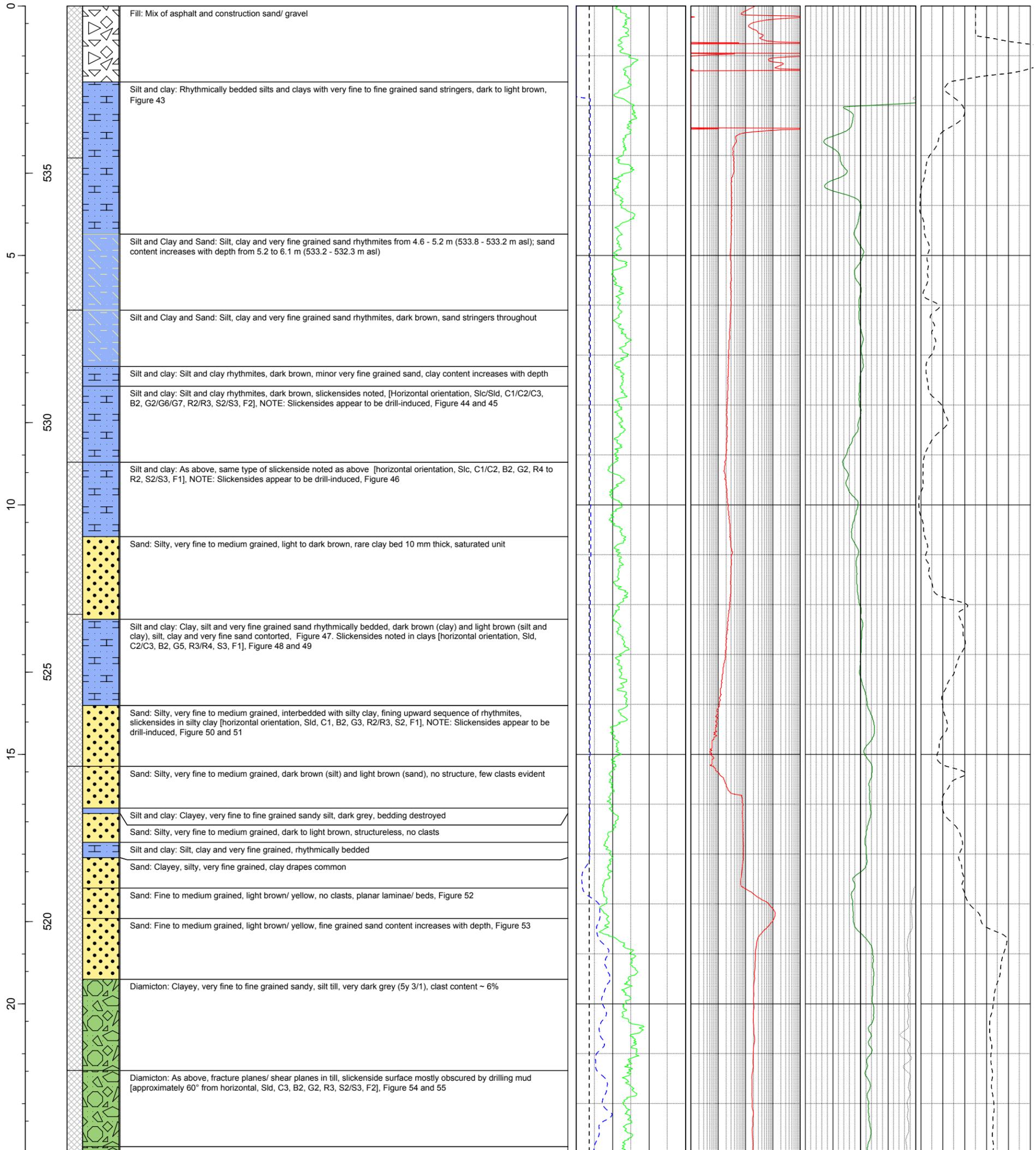
Lithology



Log ID: **PR08-5 Composite Log**

Total Depth: **177 m**  
 Location: **8-20-083-21W5M**  
 Northing UTM Nad83: **6229419**  
 Easting UTM Nad83: **483199**  
 Hole Diameter: **155 mm**  
 Elevation (Ground Surface): **538.35 m asl**  
 Drilling Date: **12/04/08**  
 Drilled By: **Boart Longyear (Sonic)/ Garrity & Baker (Wet Rotary)**  
 Lithology Logged By: **S. Slattery/ J. Morgan/ J. Warren**  
 Geophysical Log Operator: **Century Wireline Services/ D. Sinkwich**

Depth (Meters) Elevation (m asl)	Sample Interval	Lithology	Caliper (mm) 125 — 375	Resistivity (ohm.m) 0.1 — 1000	Density (CDL) (g/cc) 1 — 3	SP (mV) 100 — -100
			Bit Size (mm) 125 — 375		Comp (g/cc) 0.45 — -0.05	
			Gamma (API) 0 — 150			





Depth (Meters)  
Elevation (m asl)

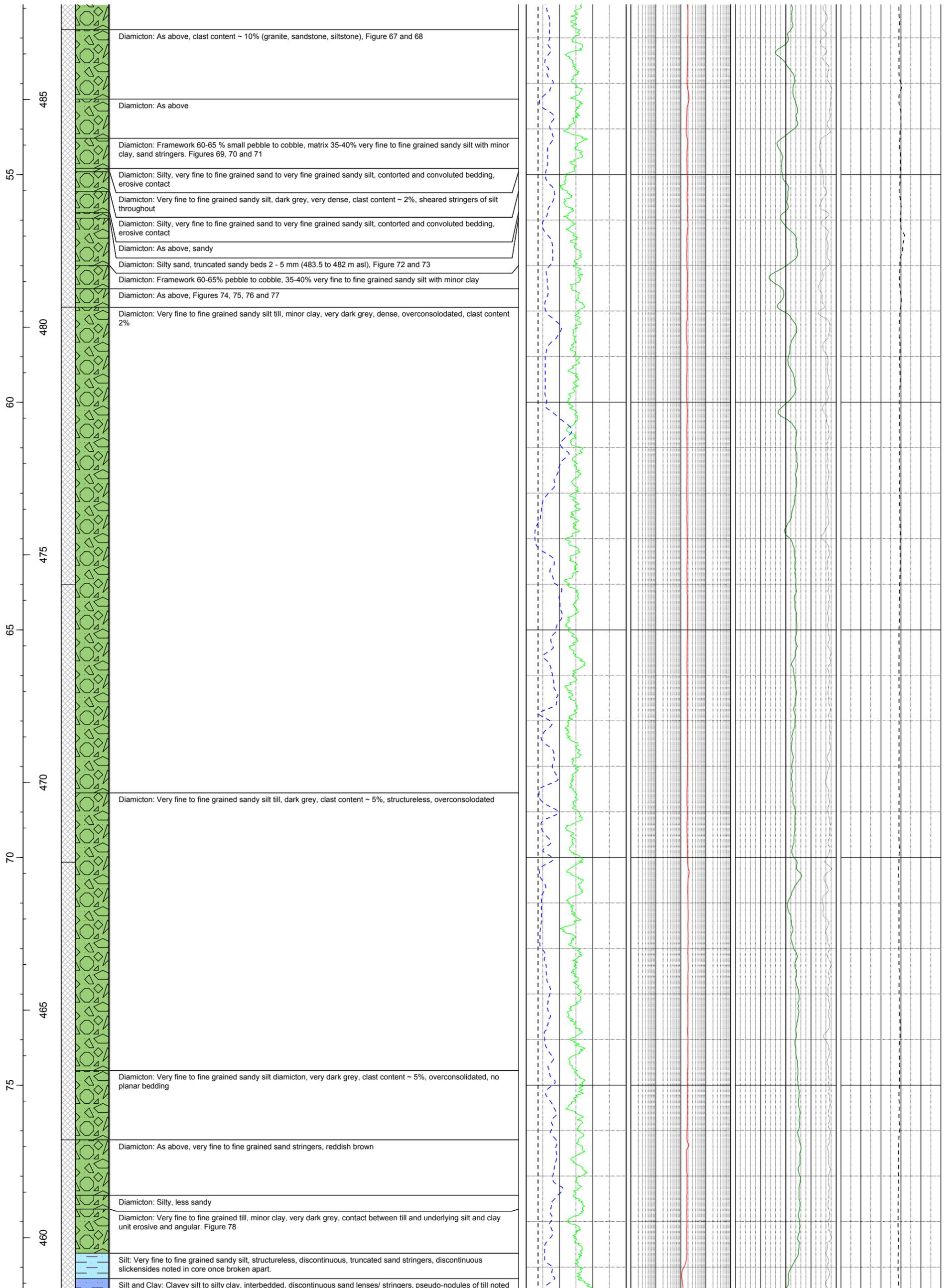
Sample Interval	Lithology
-----------------	-----------

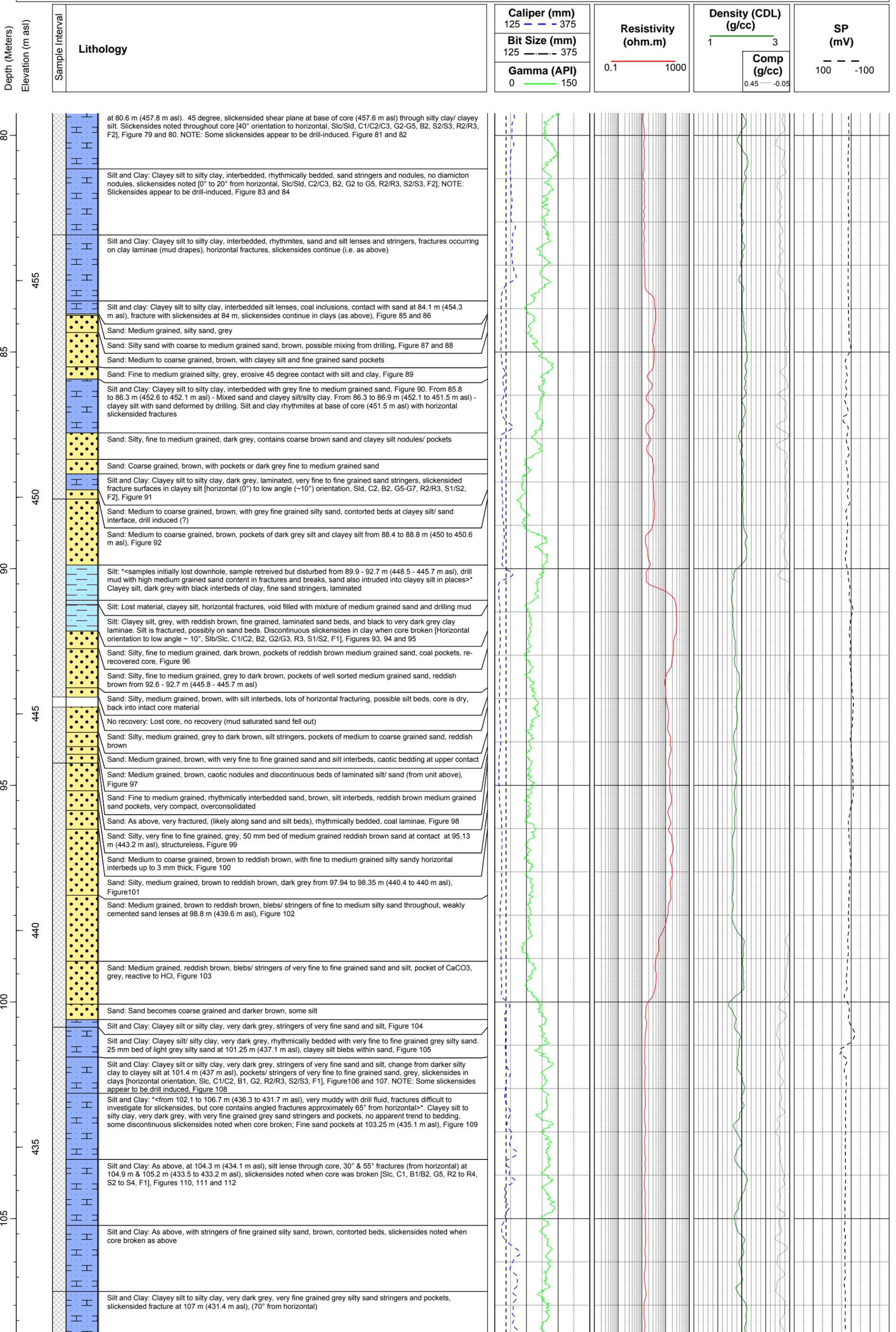
<b>Caliper (mm)</b> 125 - - - 375
<b>Bit Size (mm)</b> 125 - - - 375
<b>Gamma (API)</b> 0 - - - 150

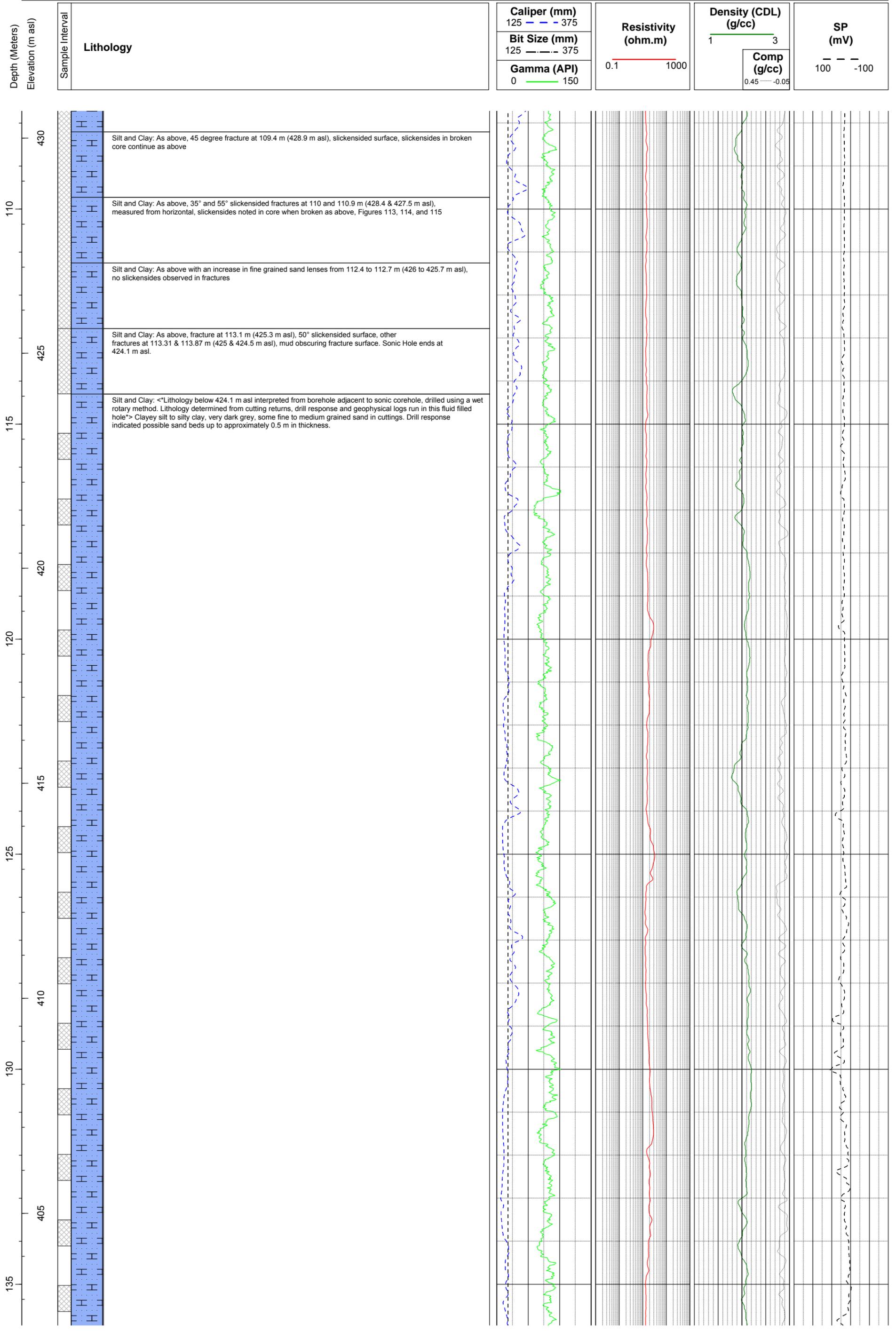
<b>Resistivity (ohm.m)</b> 0.1 - - - 1000
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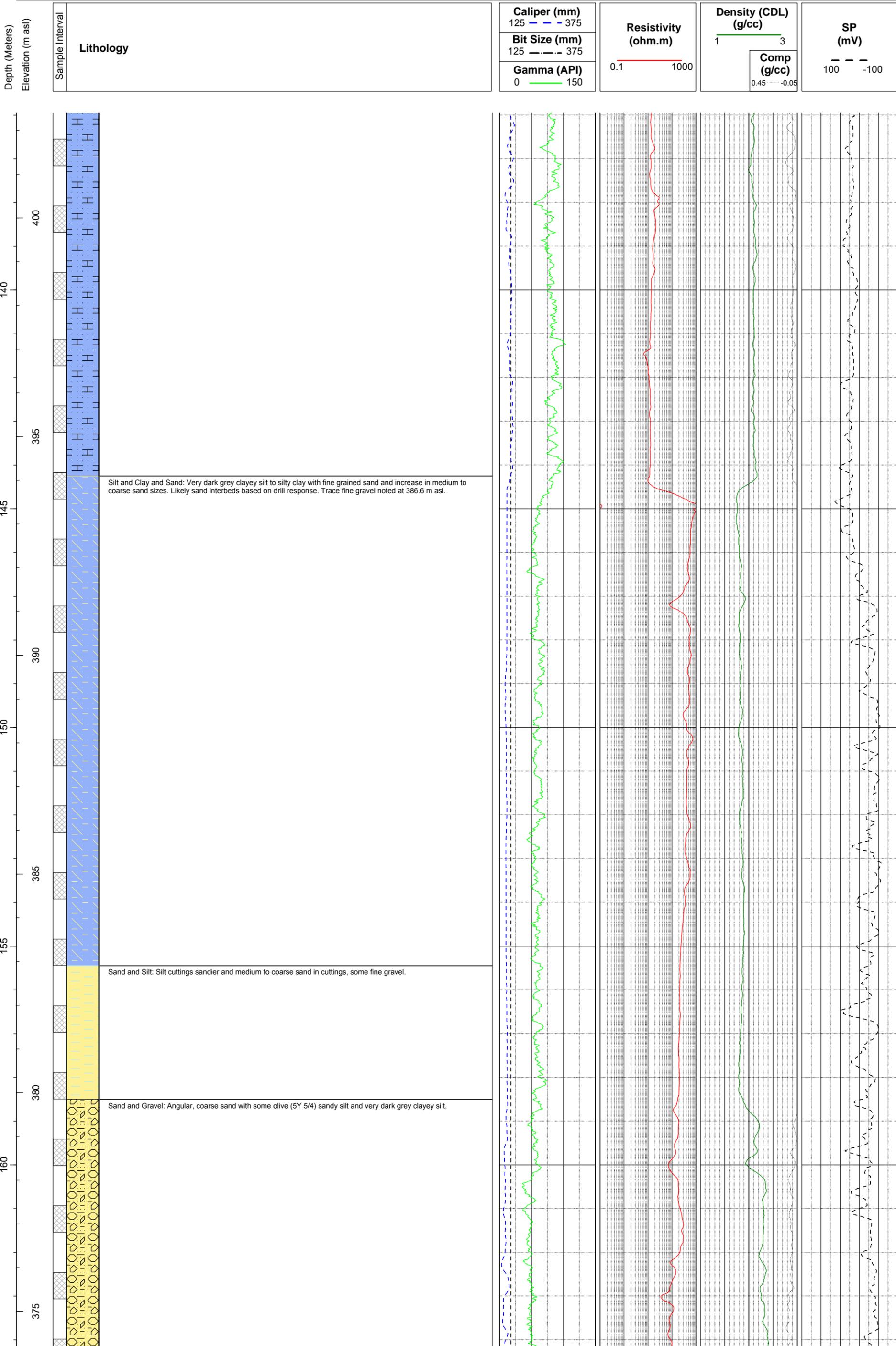
<b>Density (CDL) (g/cc)</b> 1 - - - 3
<b>Comp (g/cc)</b> 0.45 - - - 0.05

<b>SP (mV)</b> 100 - - - -100
----------------------------------









Depth (Meters)  
Elevation (m asl)

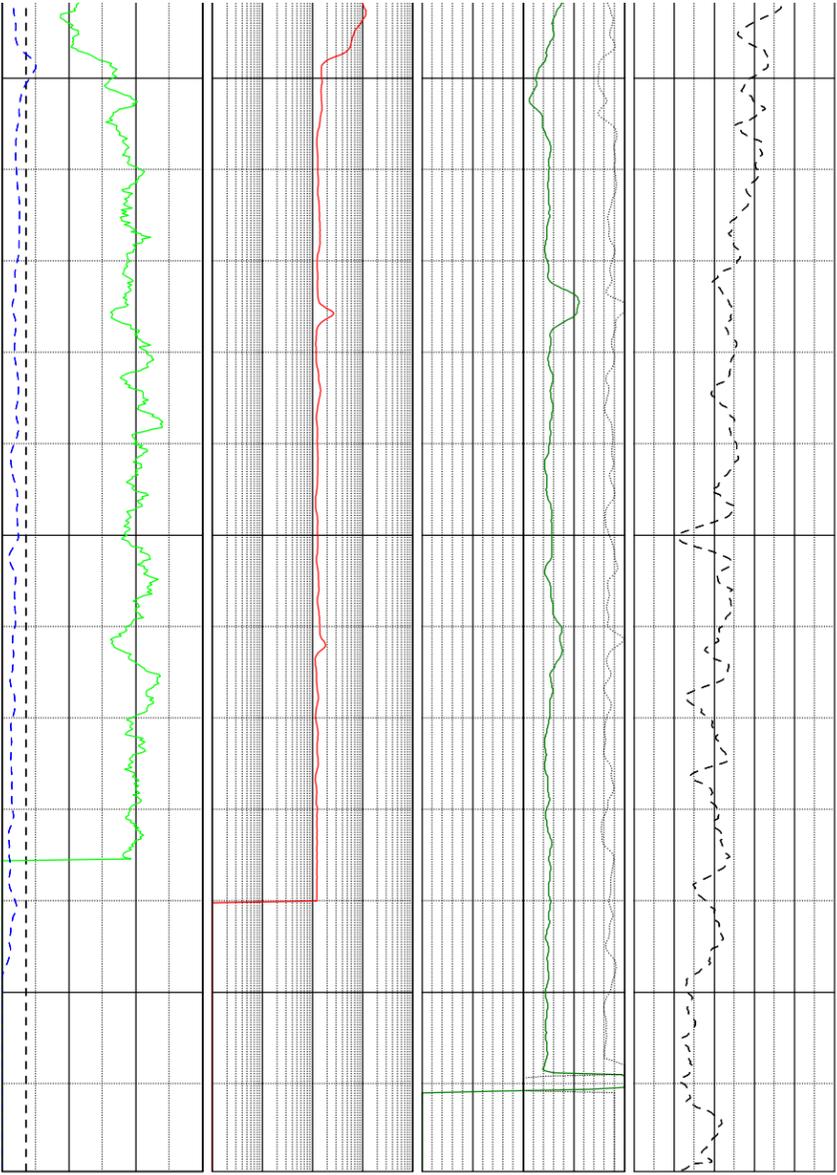
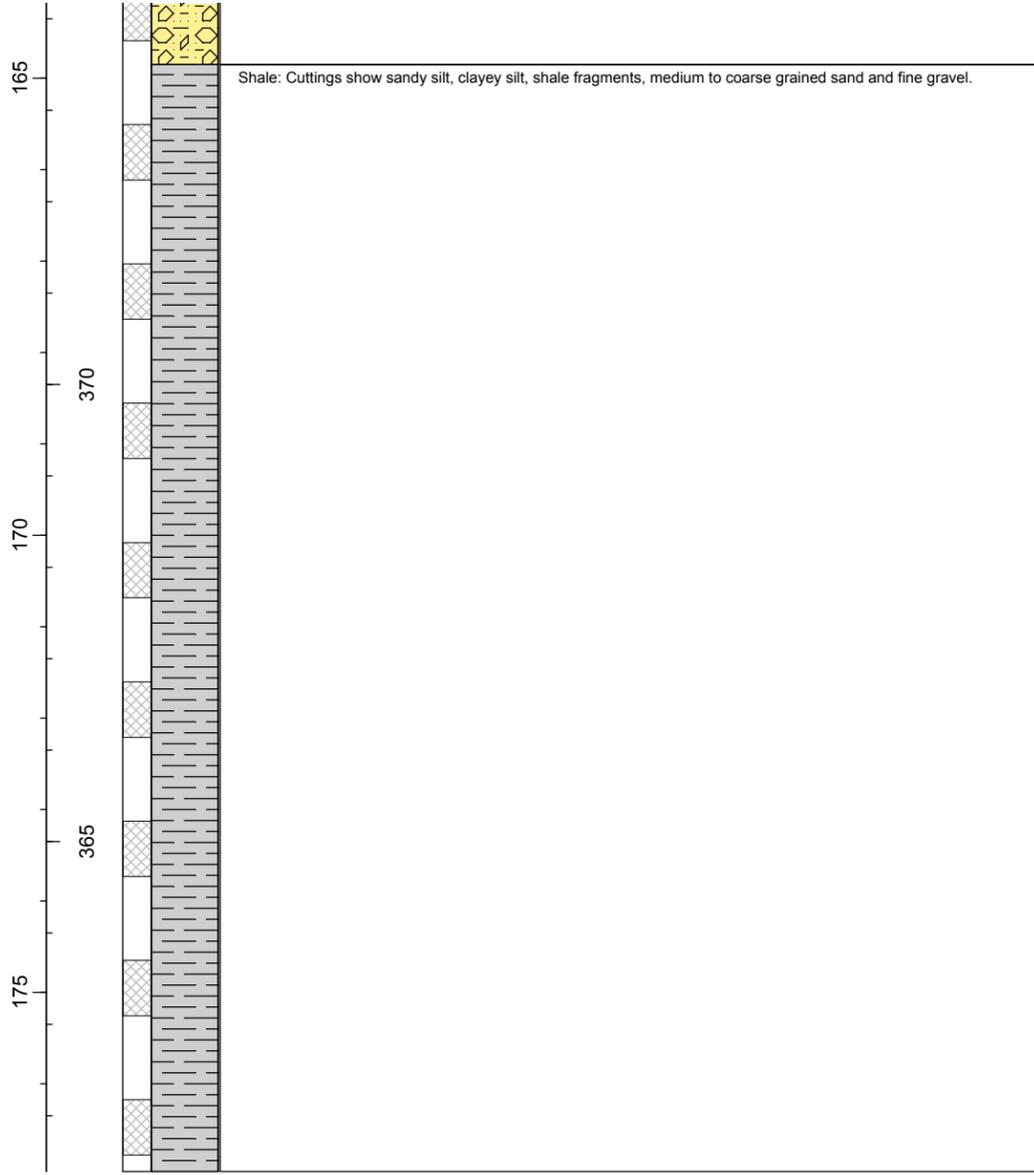
Sample Interval	Lithology
-----------------	-----------

<b>Caliper (mm)</b> 125 - - - 375
<b>Bit Size (mm)</b> 125 - - - 375
<b>Gamma (API)</b> 0 - - - 150

<b>Resistivity (ohm.m)</b>
0.1 - - - 1000

<b>Density (CDL) (g/cc)</b>
1 - - - 3
<b>Comp (g/cc)</b>
0.45 - - - 0.05

<b>SP (mV)</b>
100 - - - -100



## **Appendix 2 — Core Photographs**

Photos from borehole PR08-3 (Figures 3–42; pages 32–51)

Photos from borehole PR08-5 (Figures 43–115; pages 52–87)



Figure 3. Rhythmically bedded silt and clay (PR08-3; 13.7–14.9 m; 431.6–430.4 m asl).



Figure 4. Silty, fine- to medium-grained sand with silty, clay interclasts; (PR08-3; 15.2–16.8 m; 430.1–428.5 m asl).



Figure 5. Rhythmically bedded silt and clay (PR08-3; 18.3–19 m; 427–426.3 m asl).



Figure 6. Slickenside in silt and clay (PR08-3; 19.6 m; 425.7 m asl).



Figure 7. Clayey sandy silt to silty sand, stoney diamicton (PR08-3; 24.4–25.9 m; 420.9–419.4 m asl).



Figure 8. Slightly silty, fine- to medium-grained sand (PR08-3; 26.5–26.8 m; 418.8–418.5 m asl).



Figure 9. Interbedded sand and silt (PR08-3; 48.9–50.3 m; 396.4–395 m asl).



Figure 10. Gradational contact between (left of coin) clayey sandy silt to silty sand, stoney diamicton and (right of coin) interbedded sand and silt (PR08-3; 48.9–50.3 m; 396.4–395 m asl).



Figure 11. Deformed beds of sand and silt (PR08-3; 48.9–50.3 m; 396.4–395 m asl).



Figure 12. Clayey sandy silt to silty sand, stoney diamicton with sand lenses and blebs (PR08-3; 61–62.5 m; 384.3–382.8 m asl).



Figure 13. Clayey sandy silt to silty sand, stoney diamicton with sand lenses and blebs (PR08-3; 61–62.5 m; 384.3–382.8 masl).



Figure 14. Clayey sandy silt to silty sand, stoney diamicton with sand lenses; note shear planes through sand lenses (PR08-3; 64–65.5 m; 381.3–379.8 m asl).



Figure 15. Slickenside in clay-rich sediments within silty diamicton (PR08-3; 64 m; 381.3 m asl).



Figure 16. Clayey sandy silt to silty sand, stoney diamicton with deformed interbeds of silty sand; note ripple-drift cross stratification in sand (PR08-3; 70.1–71.7 m; 375.2–373.7 m asl).



Figure 17. Brecciated, clayey sandy silt to silty sand, stoney diamicton (PR08-3; 73.2–79.2 m; 372.2–366.1 m asl).



Figure 18. Clayey sandy silt to silty sand, stoney diamicton; slickenside in silt-rich sediments (PR08-3; 80.5 m; 364.8 m asl).



Figure 19. Clayey sandy silt to silty sand, stoney diamicton; slickensided fracture in silt-rich sediments is obscured by drilling mud (PR08-3; 84.7 m; 360.6 m asl).



Figure 20. Clayey sandy silt to silty sand, stoney diamicton; slickensided fracture in silt-rich sediments is obscured by drilling mud (PR08-3; 85.8 m; 359.5 m asl).



Figure 21. Clayey sandy silt to silty sand, stoney diamicton; slickensided fracture in silt-rich sediments is obscured by drilling mud (PR08-3; 86 m; 359.3 m asl).

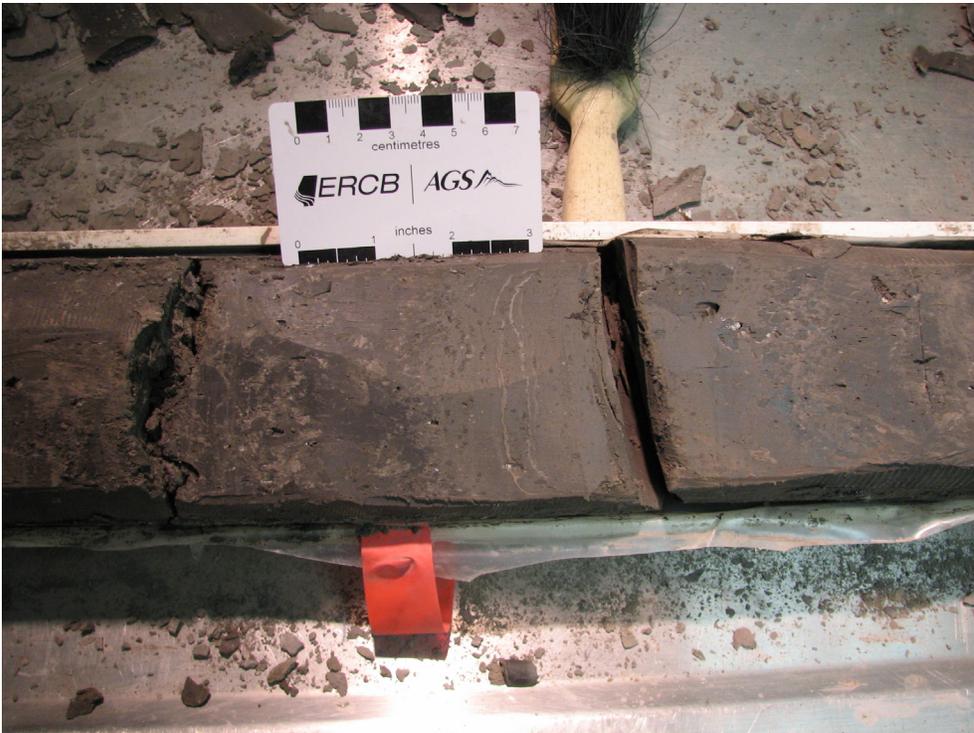


Figure 22. Rhythmically bedded deposits of very fine sandy silt and silty clay pass into brecciated deposits of silt and clay (PR08-3; 89.9 m; 355.4 m asl).



Figure 23. Rhythmically bedded deposits of very fine sandy silt and silty clay (PR08-3; 89.9 m; 355.4 m asl).



Figure 24. Deformed, rhythmically bedded deposits of very fine sandy silt and silty clay (PR08-3; 89.9–93 m; 355.4–352.3 m asl).



Figure 25. Deformed, rhythmically bedded deposits of very fine sandy silt and silty clay; discontinuous slickensides noted throughout (PR08-3; 89.9–93 m; 355.4–352.3 m asl).



Figure 26. Rhythmically bedded deposits of very fine sandy silt and silty clay; discontinuous slickensides noted throughout (PR08-3; 89.9–93 m; 355.4–352.3 m asl).



Figure 27. Rhythmically bedded deposits of very fine sandy silt and silty clay; discontinuous slickensides noted throughout (PR08-3; 89.9–93 m; 355.4–352.3 m asl).



Figure 28. Deformed, rhythmically bedded deposits of very fine sandy silt and silty clay; discontinuous slickensides noted throughout (PR08-3; 93–94.5 m; 352.3–350.8 m asl).



Figure 29. Rhythmically bedded deposits of very fine sandy silt and silty clay; discontinuous slickensides noted throughout (PR08-3; 93–94.5 m; 352.3–350.8 m asl).



Figure 30. Rhythmically bedded deposits of very fine sandy silt and silty clay (PR08-3; 94.5–96 m; 350.8–349.3 m asl).



Figure 31. Rhythmically bedded deposits of very fine sandy silt and silty clay; slickensides noted in silty clay (PR08-3; 99.1–100.6 m; 346.2–344.7 m asl).



Figure 32. Rhythmically bedded deposits of very fine sandy silt and silty clay; slickensides noted in silty clay (PR08-3; 99.1–100.6 m; 346.2–344.7 m asl).



Figure 33. Rhythmically bedded deposits of very fine sandy silt and silty clay; slickensides noted in silty clay (PR08-3; 99.1–100.6 m; 346.2–344.7 m asl).



Figure 34. Rhythmically bedded deposits of very fine sandy silt and silty clay; slickensides noted in silty clay (PR08-3; 99.1–100.6 m; 346.2–344.7 m asl).



Figure 35. Planar-bedded, silty, fine- to medium-grained sand (PR08-3; 101.6–101.7 m; 343.7–343.6 m asl).



Figure 36. Rhythmically bedded silty sand, silt and clay (PR08-3; 103.4 m; 341.9 m asl).



Figure 37. Slickensides parallel to bedding in silty clay to clayey silt (Note: slickensides appear to be drill induced), (PR08-3; 102.7–103.4 m; 342.6–341.9 m asl).



Figure 38. Slickensides parallel to bedding in silty clay to clayey silt (Note: slickensides appear to be drill-induced), (PR08-3; 103.6–105.2 m; 341.7–340.1 m asl).



Figure 39. Slickensides parallel to bedding in rhythmically bedded deposits of silty clay to clayey silt (Note: slickensides appear to be drill-induced), (PR08-3; 103.6–105.2 m; 341.7–340.1 m asl).



Figure 40. Slickensides parallel to bedding in rhythmically bedded deposits of silty clay to clayey silt (Note: slickensides appear to be drill-induced), (PR08-3; 103.6–105.2 m; 341.7–340.1 m asl).



Figure 41. Rhythmically bedded deposits of silty clay to clayey silt; note silt-rich interclasts adjacent to coin (PR08-3; 105.3–106.7 m; 340–338.6 m asl).



Figure 42. Slickensides parallel to bedding in rhythmically bedded deposits of silty clay to clayey silt (Note: slickensides appear to be drill-induced), (PR08-3; 105.3–106.7 m; 340–338.6 m asl).



Figure 43. Rhythmically bedded deposits of silt and clay (PR08-5; 1.5–4.6 m; 536.8–535.3 m asl).



Figure 44. Slickensides within silt and clay rhythmities (Note: Slickensides appear to be drill-induced), (PR08-5; 7.6–9.1 m; 530.7–529.2 m asl).



Figure 45. Slickensides within silt and clay rhythmites (Note: Slickensides appear to be drill-induced), (PR08-5; 7.6–9.1 m; 530.7–529.2 m asl).



Figure 46. Slickensides within silt and clay rhythmite (Note: Slickensides appear to be drill-induced), (PR08-5; 9.1–10.6 m; 529.2–527.7 m asl).



Figure 47. Deformed, rhythmically bedded deposits of silty clay to clayey silt (PR08-5; 12.3–14 m; 526.1–524.3 m asl).



Figure 48. Deformed, rhythmically bedded deposits of silty clay to clayey silt (PR08-5; 12.3–14 m; 526.1–524.3 m asl).



Figure 49. Slickenside in slit and clay rhythmite (PR08-5; 12.3–14 m; 526.1–524.3 m asl).



Figure 50. Silty sand with interbeds of silty clay (PR08-5; 14–15.2 m; 524.3–523.1 m asl).



Figure 51. Slickenside in silty clay (Note: Slickensides appear to be drill-induced), (PR08-5; 14–15.2 m; 524.3–523.1 m asl).



Figure 52. Fine- to medium-grained sand (PR08-5; 17.7–18.3 m; 520.7–520.1 m asl).



Figure 53. Fine- to medium-grained sand (PR08-5; 18.3–19.5 m; 520.1–518.8 m asl).



Figure 54. Fractures (jointing?) in silt-rich diamicton (PR08-5; 21.3–22.9 m; 517–515.5 m asl).



Figure 55. Fractures (jointing?) in silt-rich diamicton; surface mostly obscured by drilling mud (PR08-5; 21.3–22.9 m; 517–515.5 m asl).



Figure 56. Clayey sandy silt to silty sand, stony diamicton (PR08-5; 24.4–27.1 m; 514–511.2 m asl).



Figure 57. Erosional contact between silty sand (right of flagging tape) and overlying clayey sandy silt to silty sand, stony diamicton (left of flagging tape; 28.8–29 m; 509.6–509.4 m asl).

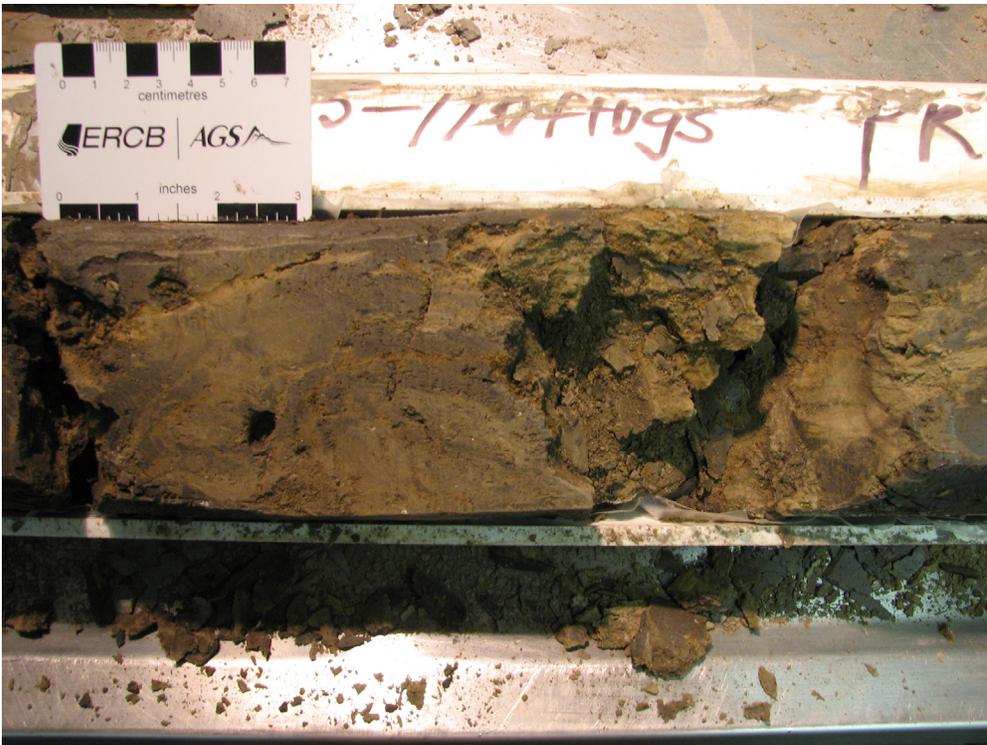


Figure 58. Fine- to medium-grained sand with interbeds of diamicton; ripple drift in sand, contorted and overturned (PR08-5; 32.5–33.2 m; 505.8–505.1 m asl).



Figure 59. Fine- to medium-grained sand with interbeds of till; ripple drift in sand, contorted and overturned; contact with diamicton shown (PR08-5; 32.5–33.2 m; 505.8–505.1 m asl).



Figure 60. Interclasts of very fine grained sand (PR08-5; right of scale card) in clayey sandy silt to silty sand, stony diamicton (PR08-5; 44.2–45.7 m; 494.2–492.6 m asl).



Figure 61. Beds of silty, very fine to fine-grained sand in clayey sandy silt to silty sand, stony diamicton (PR08-5; 44.2–45.7 m; 494.2–492.6 m asl).



Figure 62. Planar bedded deposits of silty fine sand and silt (PR08-5; 45.7–46.3 m; 492.6–492 m asl).



Figure 63. Clayey sandy silt to silty sand, stony diamicton (PR08-5; 46.3–47.2 m; 492–491 m asl).



Figure 64. Contorted beds of sandy silt and silty sand (PR08-5; 47.2–48.8 m; 491.1–489.6 m asl).



Figure 65. Beds of sandy silt and silty sand, blocks of dislocated clayey sandy silt to silty sand, stony diamicton (PR08-5; 48.8–50.3; 489.6–488.1 m asl).



Figure 66. Contact between sandy silt and clayey sandy silt to silty sand, stony diamicton; note load structure immediately below coin (PR08-5; 50.3–50.8 m; 488.1–487.6 m asl).





Figure 69. Clayey sandy silt to silty sand, stony diamicton; clast size ranges between small pebble and cobble grade (PR08-5; 54.2–54.9 m; 484.2–483.5 m asl).



Figure 70. Clayey sandy silt to silty sand, stony diamicton; clast size ranges between small pebble and cobble grade (PR08-5; 54.2–54.9 m; 484.2–483.5 m asl).



Figure 71. Clayey sandy silt to silty sand, stony diamict; clast size ranges between small pebble and cobble grade (PR08-5; 54.2–54.9 m; 484.2–483.5 m asl).

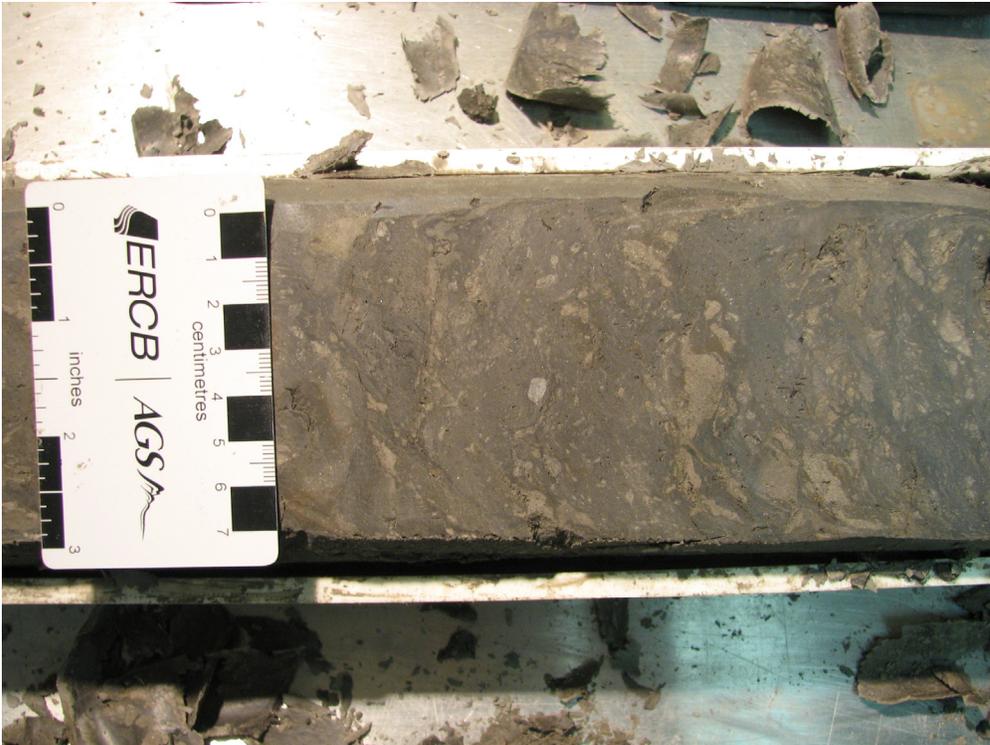


Figure 72. Clayey sandy silt to silty sand, stone-poor diamict (PR08-5; 56–56.4 m; 482.4–481.9 m asl).



Figure 73. Clayey sandy silt to silty sand, stony diamicton (PR08-5; 56–56.4 m; 482.4–481.9 m asl).



Figure 74. Clayey sandy silt to silty sand, stony diamicton; clast size ranges between small pebble and cobble grade; note cobble lag below scale card (PR08-5; 56.4–57.9 m; 482.4–481.9 m asl).



Figure 75. Clayey sandy silt to silty sand, stony diamicton; clast size ranges between small pebble and cobble grade; note cobble lag below scale card (PR08-5; 56.4–57.9 m; 482.4–481.9 m asl).



Figure 76. Clayey sandy silt to silty sand, stony diamicton; clast size ranges between small pebble and cobble grade (PR08-5; 56.4–57.9 m; 482.4–481.9 m asl).



Figure 77. Clayey sandy silt to silty sand, stony diamicton; note interclasts of silty sand to sandy silt (PR08-5; 56.4–57.9 m; 482.4–481.9 m asl).



Figure 78. Erosive contact between clayey sandy silt to silty sand, stony diamicton and underlying silt and clay (PR08-5; 77.7–78.7 m; 460.6–459.7 m asl).



Figure 79. 45° slickenside in silty clay to clayey silt (PR08-5; 80.8 m; 457.6 m asl).



Figure 80. 45° slickenside in silty clay to clayey silt (PR08-5; 80.8 m; 457.6 m asl).



Figure 81. Slickensides in silty clay to clayey silt (Note: Slickensides appear to be drill-induced), (PR08-5; 79.2–80.8 m; 459.1–457.6 m asl).



Figure 82. Slickensides in silty clay to clayey silt (Note: Slickensides appear to be drill-induced), (PR08-5; 79.2–80.8 m; 459.1–457.6 m asl).



Figure 83. Slickensides within clayey silt to silty clay (Note: Slickensides appear to be drill-induced), (PR08-5; 80.8–82.3 m; 457.6–456.1 m asl).



Figure 84. Slickensides within clayey silt to silty clay (Note: Slickensides appear to be drill-induced), (PR08-5; 80.8–82.3 m; 457.6–456.1 m asl).



Figure 85. Gradational contact between silty clay to clayey silt and fine- to medium-grained sand (PR08-5; 84.1 m; 454.3 m asl).



Figure 86. Slickenside within silty clay to clayey silt (PR08-5; 83.8–84.1 m; 454.5–454.2 m asl).



Figure 87. Medium- to coarse-grained silty sand with interclasts of clayey silt (PR08-5; 84.2–84.6 m; 454.2–453.8 m asl).



Figure 88. Medium- to coarse-grained silty sand with interclasts of clayey silt (PR08-5; 84.2–84.6 m; 454.2–453.8 m asl).



Figure 89. Erosive contact between fine- to medium-grained silty sand and silty clay to clayey silt (PR08-5; 85.3–85.6 m; 453–452.7 m asl).



Figure 90. Clayey silt to silty clay interbedded with silty, fine- to medium-grained sand (PR08-5; 85.6–86.9 m; 452.7–451.5 m asl).



Figure 91. Clayey silt to silty clay, laminated, with horizontal slickenside surfaces (PR08-5; 87.8–88.2 m; 450.5–450.1 m asl).



Figure 92. Medium- to coarse-grained sand (PR08-5; 88.4–89.9 m; 450–448.4 m asl).



Figure 93. Rhythmically bedded deposits of clayey silt, silty clay and very fine to fine-grained sand; sample lost downhole and re-recovered (PR08-5; 90.8–91.4 m; 447.5–446.9 m asl).



Figure 94. Rhythmically bedded deposits of clayey silt, silty clay and very fine to fine-grained sand; slickensides in clay-rich sediments; sample lost downhole and re-recovered (PR08-5; 90.8–91.4 m; 447.5–446.9 m asl).



Figure 95. Rhythmically bedded deposits of clayey silt, silty clay and very fine to fine-grained sand; slickensides in clay-rich sediments; sample lost downhole and re-recovered (PR08-5; 90.8–91.4 m; 447.5–446.9 m asl).



Figure 96. Silty, fine- to medium-grained sand; sample lost downhole and re-recovered (PR08-5; 91.4–91.8 m; 446.9–446.5 m asl).



Figure 97. Fine- to medium-grained sand, discontinuous beds (laminae) of silty sand to sandy silt (PR08-5; 94.1–94.3 m; 444.2–444.1 m asl).



Figure 98. Fine- to medium-grained sand with silty interbeds, weakly cemented and fractured (PR08-5; 94.5–95.1 m; 443.9–443.2 m asl).



Figure 99. Very fine to fine-grained silty sand (PR08-5; 95.1–95.6 m; 443.2–442.8 m asl).



Figure 100. Medium- to coarse-grained sand with interbeds of silt (PR08-5; 95.6–96 m; 442.8–442.3 m asl).



Figure 101. Medium-grained silty sand (PR08-5; 96–97.5 m; 442.3–440.8 m asl).

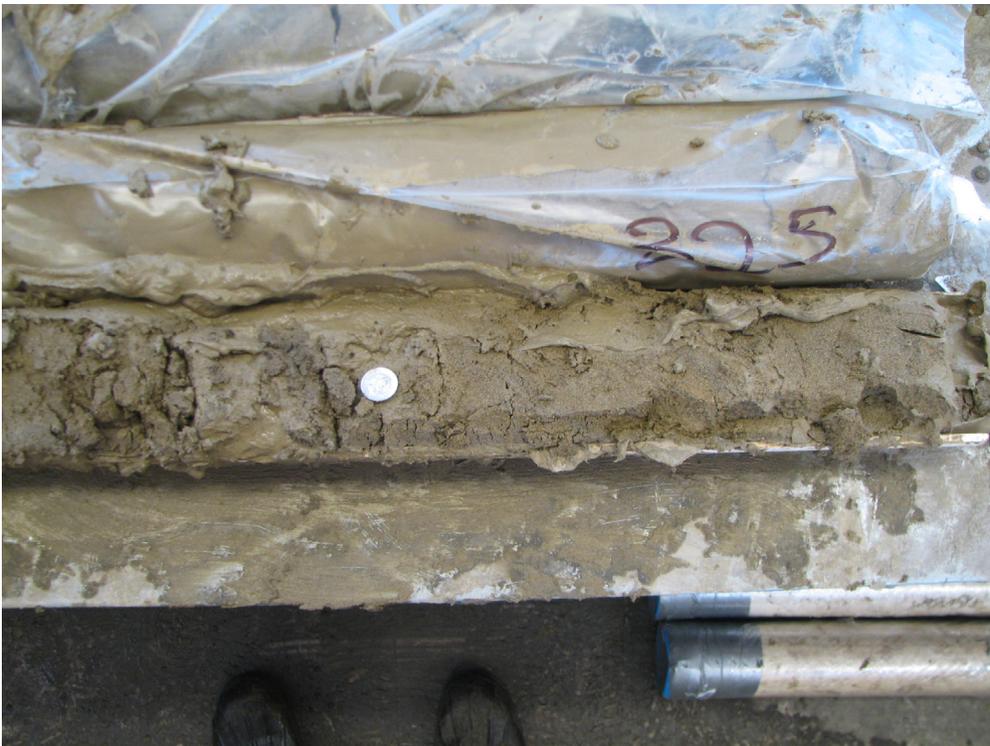


Figure 102. Medium-grained sand with blebs of silty sand (PR08-5; 97.5–99.1 m; 440.8–439.3 m asl).



Figure 103. Carbonate interclast (right of coin) in medium-grained sands (PR08-5; 99.1–100.1 m; 439.3–438.3 m asl).



Figure 104. Erosive contact between silty fine sand and silty clay (PR08-5; 100.4–100.6 m; 438–437.8 m asl).



Figure 105. Clayey silt to silty clay interbedded with silty sand; note 2.5 cm thick bed of silty sand (PR08-5; 100.6–101.3 m; 437.8–437.1 m asl).



Figure 106. Clayey silt to silty clay, wisps of silty, very fine sands throughout (PR08-5; 101.3–102.1 m; 437.1–436.2 m asl).



Figure 107. Clayey silt to silty clay, wisps of silty, very fine sand throughout (PR08-5; 101.3–102.1 m; 437.1–436.2 m asl).



Figure 108. Clayey silt to silty clay, wisps of silty, very fine sand throughout; slickensides in clay (Note: Slickensides appear to be drill-induced), (PR08-5; 101.3–102.1 m; 437.1–436.2 m asl).



Figure 109. Clayey silt to silty clay, note wisps of silty, very fine sands (PR08-5; 102.1–103.6 m; 436.2–434.7 m asl).



Figure 110. Silt-rich lens in clayey silt to silty clay (PR08-5; 104.3 m; 434.1 m asl).



Figure 111. Slickensides in silty clay to clayey silt (PR08-5; 104.9–105.2 m; 433.5–433.2 m asl).



Figure 112. Slickensides in silty clay to clayey silt (PR08-5; 104.9–105.2 m; 433.5–433.2 m asl).



Figure 113. Silt and clay with very fine grained sand stringers and blebs (PR08-5; interclasts; 110–110.9 m; 428.4–427.5 m asl).



Figure 114. Slickensides in silt and clay (PR08-5; 110–110.9 m; 428.4–427.5 m asl).



Figure 115. Slickensides in silt and clay (PR08-5; 110–110.9 m; 428.4–427.5 m asl).

## Appendix 3 — Geotechnical Logging of Discontinuities in Stiff Clays and Argillaceous Bedrock

### Introduction

To provide a realistic range of strength parameters for materials, thorough geotechnical logging of core is required to identify the frequency, orientation and condition of all potential structural defects or discontinuities. This document outlines a classification system developed to aid in describing discontinuities in a repeatable manner, such that end users have a clear concept of their potential effects on the strength or behaviour of the material.

For more than 30 years, ‘hard-rock’ geology has used repeatable systems to classify discontinuities. When such systems were applied to Quaternary sediments and weakly cemented rocks, such as Cretaceous shales, their limitations became apparent because sediments and soft rocks deform in a more ductile manner than ‘hard’ rocks. This ductile deformation leads to markedly different shapes and conditions of discontinuities and shear surfaces in these materials.

The main causes of shear surfaces in Quaternary deposits and the immediately underlying bedrock in northern Alberta include

- glacial shear/thrust,
- stress rebound,
- differential compaction, and
- large-displacement shears.

Based on the differences in the deformation properties of soft rocks, we undertook a literature review for classification of discontinuities in overconsolidated sediments and soft rocks. This encountered work undertaken at Imperial College in the late 1960s by Fookes (Fookes and Denness, 1969).

Based on this information and experience with coreholes logged for geotechnical design purposes, we developed a classification system derived from work by Fookes (United States Department of the Interior, 1998), as well as classifications encountered during experience logging Cretaceous shales in the consulting industry.

The purpose of this system is to provide a means by which geologists, technologists and engineers can log core in a repeatable manner that will give end users an increased level of confidence in the choice of soil or rock parameters and/or behavioural characteristics.

### Nomenclature

#### *Spacing*

The quantification of the spacing of discontinuities in the core gives an indication of whether or not there is a discrete plane of weakness or deformed zone. This also helps in characterizing the potential origin of the discontinuities. The classification in Table 2 quantifies discontinuity spacing.

**Table 2. Classification of spacing.**

ID	Frequency	Spacing
Slx	Extensive	Completely slickensided (no bedding structure)
Sla	Abundant	<1 cm
Slb	Numerous	1–3 cm
Slc	Occasional	3–6 cm
Sld	Isolated	>6 cm
Slo	Slo	None

### **Continuity**

A continuous joint or fracture is weaker and more deformable than a short, discontinuous fracture bridged by intact bedrock. Identification of the more continuous fractures is an important aspect of formulating rock stability input data. Table 3 provides descriptors for continuity.

**Table 3. Classification of continuity.**

ID	Continuity	Description
C1	Discontinuous	<50% through sample
C2	Moderately continuous	50%–95% through sample
C3	Very continuous	Completely through sample

### **Core Breakage**

Description of the physical force required to break the discontinuities (Table 4) is a crude but effective method of describing the degree to which a discontinuity has healed. If it takes physical force to break, the discontinuity is likely well healed, which may allow for a more optimistic estimate of the shear strength.

**Table 4. Classification of cohesiveness.**

ID	Force Applied	Result
B1	Mechanical force	Requires force to break
B2	Hand force	Breaks apart in hand
B3	No force	Falls apart

### **Geometry**

The geometry of the discontinuity can provide important information regarding the origin of the discontinuity. It is most common that planar features are associated with bedding or large displacement shears and that many of the irregular surfaces are the product of consolidation or weathering. Table 5 provides a list of descriptors and schematics for geometry. See Fookes and Denness (1969) for a more detailed explanation.

**Table 5. Classification of geometry.**

ID	Geometry
G1	Planar
G2	Semi-curved
G3	Curved
G4	Hinged
G5	Semi-undulose
G6	Undulose
G7	Conchoidal

**Roughness**

The roughness (small-scale asperities) of discontinuity surfaces is critical for evaluating shear strength. Roughness descriptors should be used whenever they are observed. Table 6 provides descriptions of degrees of roughness. See Fookes and Denness (1969) for a more detailed explanation, as well as visual examples of degrees of roughness.

**Table 6. Classification of roughness.**

ID	Roughness
R1	Striated
R2	Very smooth (polished)
R3	Smooth
R4	Slightly rough
R5	Very rough
R6	Pock Marked
R7	Pitted

**Sheen**

The sheen describes the degree to which the surface of the discontinuity has been sheared and reworked. A glossy surface likely has the majority of surface defects polished off and is likely closer to a full residual strength. Table 7 provides descriptors for describing sheen.

**Table 7. Classification of sheen.**

ID	Sheen
S1	Very high gloss (mirror like)
S2	Shiny, glossy
S3	Moderately shiny
S4	Dull

## Orientation

The orientation of all fracture with respect to applied loads can be critical to deformation or stability. It is important to distinguish between discontinuities and shear surfaces that form along bedding and those that extend through the rock mass. Table 8 provides descriptors and Figure 116 provides schematic representations of discontinuity orientation.

Table 8. Classification of orientation.

ID	Orientation	Description
F1	Unidirectional (parallel)	Parallel to bedding
F2	Unidirectional (non parallel)	Not parallel to bedding
F2	Multidirectional	In multiple directions
F3	Intersecting	Planes intersect one another

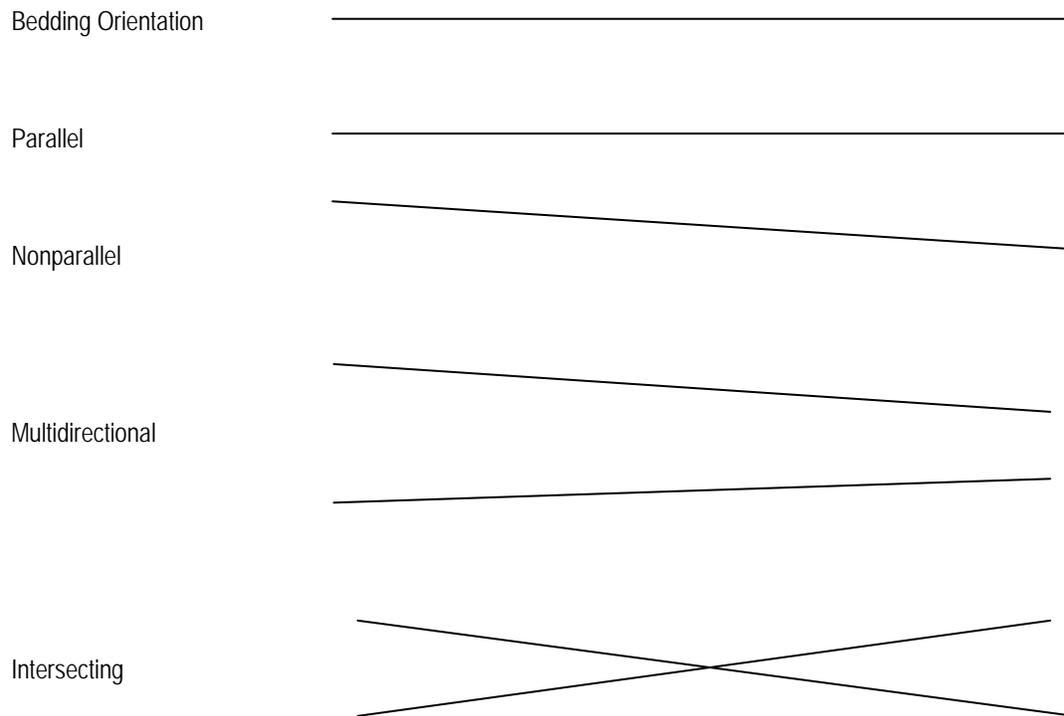


Figure 116. Schematic representations of discontinuity orientation.

## Examples of Logged Samples



Figure 117. Discontinuity example: continuous, planar, moderately shiny, striated. Scale indicated by gloved hand holding the 7.6 cm (3 inch) diameter core sample.



Figure 118. Discontinuity example: continuous, planar, slightly rough, dull. Scale indicated by gloved hand holding the 7.6 cm (3 inch) diameter core sample.



Figure 119. Discontinuity Example: Smooth, undulating, polished, continuous, intersecting. Scale indicated by gloved hand holding the 7.6 cm (3 inch) diameter core sample.



Figure 120. Discontinuity example: continuous, curved, striated, polished.



Figure 121. Discontinuity example: continuous, striated, slightly rough, intersecting (7.6 cm [3 inch] diameter core sample).



Figure 122. Discontinuity example: continuous, semi-undulose, striated, moderately shiny (7.6 cm [3 inch] diameter core sample).



Figure 123. Discontinuity example: continuous, semi-undulose, very shiny, striated, intersecting (7.6 cm [3 inch] diameter core sample).



Figure 124. Discontinuity example: continuous, semi-undulose, striated, polished, intersecting (7.6 cm [3 inch] diameter core sample).