

Thickness of drift in metres less than 15 >30 to 60 C.P. Kathol 1973 € 1976 L.D. Andriashek M.M. Fenton MAPPING COVERAGE Geology by: L.D. Andriashek, 1973-1976; M.M. Fenton, 1974-1975; J.D. Root, 1973. GLACIOLACUSTRINE DEPOSITS zite erratics. It can be subdivided into a clayey till (30 percent sand, S.D. = 2.5; 40 percent clay, S.D. = 3.4; N. = 36) and a sandy till (40 percent sand, S.D. = 2.6; 33 percent clay, S.D. = 3.3; N. = 24). over submerged ice in the Mayerthorpe area. Later melting of this buried areas of glacially thrust bedrock. These variations are probably the result of local incorporation of underlying bedrock. ice produced the present-day hummocky terrain. In the east a pitted delta (unit sLGp) was formed by streams flowing

BEDROCK GEOLOGY

LEGEND

PASKAPOO FORMATION: sandstone, siltstone and mudstone; minor

WHITEMUD AND BATTLE FORMATIONS: white-weathering, bentonitic

sandstone: clay and silty clay (Whitemud Formation)-purplish

Surface contour (contour interval 500 feet) ..... - 2500 -

black, mauve-weathering, bentonitic mudstone (Battle Formation)

TKp conglomerate; Scollard Member (Ksc): sandstone; bentonitic

Kwt WAPITI FORMATION: gray sandstone; mudstone and bentonite; scattered coal beds

TERTIARY AND CRETACEOUS

mudstone; thick coal beds

THALWEGS AND DRIFT THICKNESS

From R. Green, 1972

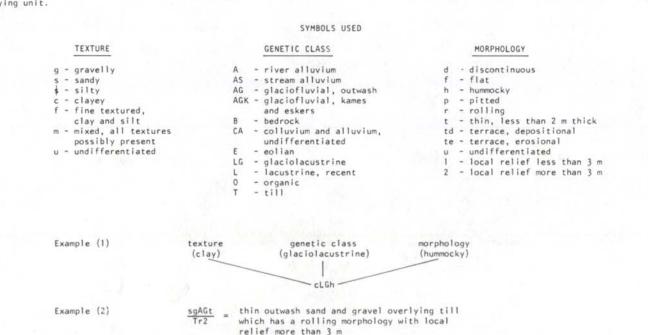
Scale 1:1,000,000

Modified from V.A. Carlson, 1970

SURFICIAL GEOLOGY WABAMUN LAKE

One term placed above another, for example  $\frac{L_0}{T}$ , indicates one deposit overlying another. The map units show the material to be expected to a depth of 2 to 3 m. The color indicates the main unit and the pattern, if present, indicates the overlying unit.

EXPLANATION OF UNIT SYMBOLS



ALBERTA

**NTS 83G** 

The mapping scheme used in this map is modified from that of the Geological Survey of Canada. The west half and northeast quarter of the area were mapped using this scheme; however, the southeast quarter was mapped following Bayrock's (1972) scheme and the information later converted to Preliminary, 1:50,000 scale, surficial geology maps, logs of 68 testholes,

and the data from analyzed samples are on open file at the Alberta Research Council.

The authors acknowledge Alberta Environment for supplying testhole logs and for discussions on the preglacial topography of the area. Physiography: The area lies within the Alberta Plains; the land slopes from an elevation of 1200 m in the southwest to 700 m in the hortheast. The west half of the area consists of broad, bedrock-controlled, ridges and valleys, formed preglacially and mantled by a generally thin (<15 m) cover of drift. In the east half the morphology is primarily the result of glacial and glaciolacustrine deposition, the former producing hummocky to rolling terrain to the north and south and the latter a flat to pitted plain in the east and south-center. Drift masks all but the general outline of preglacial valleys.

incised into bedrock or preglacial sediment. The larger lakes fill broad lows in the post-glacial landscape (Chip Lake, Pigeon Lake, Lac St. Anne); or depressions formed, and/or enlarged by glacial thrusting (Lake Wabamun, Lac La Nonne, Sandy Lake). Drift Thickness and Preglacial Topography: The preglacial land surface slopes northeastward and is cut by a number of valleys eroded during the Late Tertiary (Carlson, 1979). The drift thickness ranges from less than 1 m to more than 80 m.

SURFICIAL DEPOSITS

The North Saskatchewan, McLeod, and Pembina Rivers are generally deeply

Till is unsorted sediment deposited directly from a glacier and is composed of varying proportions of all sizes of materials eroded by the Till forms the surface deposits over more than half the area and in most places underlies the lacustrine, eolian, and organic sediment. The till is subdivided on the basis of morphology (hummocky, rolling, flat) and local relief (more than 3 m, less than 3 m) into a number of mappable units. The landform type does not imply genesis; for example, hummocky till (Th) includes both sediment deposited by stagment ice processes and that formed by other processes such as glacial thrusting. Composition: The surface till can be subdivided into materials of Laurentide provenance and mixed Laurentide and Cordilleran provenance. The "mixed till" contains a relatively large number of carbonate clasts and quartzite erratics up to 2 m<sup>3</sup>. The average texture is 40 percent sand (standard deviation = 3.2) and 25 percent clay (S.D. = 4.2; N. = 8). This till is present only in Tps. 47 and 48, Rs. 13 and 14. The boundary between this and the Laurentide till is approximate and is based on field evidence and on the soils map of Twardy and Lindsay (1971).

The Laurentide till contains few carbonate clasts and no large quart-Each subunit forms the surface till at many sites. Local variations n till texture include three sites west of R. 10 (30 percent sand, S.D. = 2.5; 31 percent clay, S.D. = 1.7), three sites west of Pigeon Lake (48 percent sand, S.D. = 4.6; 19 percent clay, S.D. = 4.5), and

Stratigraphy: The stratigraphic relationship between the two Laurentide till subunits is uncertain. The clayey till overlies the sandy till in five testholes (Tp. 52, R. 13; Tp. 48, 49, and 50, R. 9; Tp. 55, R. 2) and in the eastern one these tills are separated by stratified drift. The sandy till overlies stratified till and a clay till in two places (Tp. 54, R. 4; Tp. 57, R. 6). In a third hole the sequence is, from the top, clayey till/sandy till/sand/clayey till which suggests either (1) two clayey tills or (2) stacking of clayey till by glacial thrusting.

Glacial Thrusting: Glacial thrusting refers to the process whereby masses of underlying material are removed more-or-less intact by the glacier and The thrust sediment may include till, clay, silt, sand, gravel (Babcock, Fenton and Andriashek, 1978), and, in many areas, a large amount of bedrock material. Thrusting can extend to a depth of at least 30 m. The thrust sediment may be deposited immediately down glacier of the source depression or moved farther, and form either small hills (<2 m high,

<0.02 km²) or large hills and ridges (>10 m high, >2.5 km²). The glacially thrust terrain is extensive in the east and has been recognized in many places where the bedrock is near the surface (unit  $\frac{Tt}{R}$ ) in the area covered by lacustrine sediment southwest of Lake Wabamun, and

southwest of a line through Lac La Nonne and Sandy Lake. This terrain has been recognized in only a few areas in the west: around Chip Lake and GLACIOFLUVIAL DEPOSITS posits in and adjacent to most major meltwater channels. In the west major sources of sand and gravel exist in terraces in Tps. 46 to 50,

Rs. 12 to 14 and along the upper terraces of the McLeod, Pembina, and North Saskatchewan Rivers. In the east extensive gravel deposits are found along the meltwater channel system in the Onoway - Heatherdown -Duffield area. Near Heatherdown the terraces contain quartzite gravel reworked from the preglacial Onoway Channel. Near Duffield, late-stage erosion left only a discontinuous lag gravel, and in places exposed bedrock. Excavation into one esker located in Tp. 47, R. 13 has exposed 10 m of sand and gravel and in another (Tp. 53, R. 3), 6 m of sand beneath a till veneer. A kame moraine in Tp. 55, R. 6 contains discontinuous bodies of poorly sorted fluvial sediments and till.

approximately 40 percent of the area is covered by sediments deposited in meltwater lakes ponded against the northeast retreating glacier. In the west half these deposits, mostly well-bedded silts and clays, form a thin to discontinuous drape on the highlands and a thicker cover in the valleys. Some of the thickest sediment is massive clay deposited

southeast off the glacier and burying stagnant ice. This ice later melted to produce the pits with the deepest situated in the northwest adjacent to the glacier-delta contact. This unit grades into generally finer textured and thinner glaciolacustrine sediment (units mLG, sLG,

EOLIAN DEPOSITS Eolian deposits in the map area consist mainly of U-shaped sand dunes located on the south-east margins of major meltwater channels. All dunes are presently inactive and commonly organic deposits occupy the inter-dune areas. Individual dunes attain a height of 15 m in Tp. 47 and 56, R. 12 to 13. The deltaic deposits furnished sediment for a small dune area (Tp. 50, R. 1). Loess deposits were not recognized in

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the study area.

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