RESEARCH COUNCIL OF ALBERTA

BEDROCK GEOLOGY OF NORTHERN ALBERTA

LEGEND

CRETACEOUS

WAPITI FORMATION: grey carbonaceous feldspathic sandstone, silty shale, bentonite; thin coal beds; nonmarine

27 PUSKWASKAU FORMATION: dark grey fossiliferous shale, silty in upper

BAD HEART FORMATION: fine-grained quartzose sandstone, ferruginous oolitic sandstone and mudstone; marine

KASKAPAU FORMATION: dark grey silty shale, ironstone partings; interbedded in lower part with fine-grained quartzose sandstone and thin beds of ferruginous oolitic mudstone; marine

24b UNNAMED DARK GREY SHALE, Caribou Mountains and Buffalo Head

SMOKY GROUP: dark grey shale and silty shale, ironstone partings and concretions; marine

LABICHE FORMATION: dark grey shale and silty shale, ironstone partings and concretions; thin fish scale-bearing silty beds in lower part; marine

DUNVEGAN FORMATION: grey fine-grained feldspathic sandstone with hard calcareous beds, laminated siltstone and grey silty shale; deltaic to marine

SHAFTESBURY FORMATION: dark grey fish scale-bearing shale, silty in upper part; numerous nodules and thin beds of concretionary ironstone; bentonite partings; interbedded locally in lower part with thin silty and sandy intervals; marine

PELICAN FORMATION: fine-grained quartzose sandstone, silty and glauconitic in lower part; marine

19 JOLI FOU FORMATION: dark grey fossiliferous shale, silty interbeds in upper

PEACE RIVER FORMATION, CADOTTE MEMBER (upper): fine-grained quartzose sandstone; HARMON MEMBER (middle): dark grey silty shale; NOTIKEWIN MEMBER (lower): fine-grained glauconitic sandstone, silty interbeds in lower part; marine

ALICE CREEK TONGUE, GRAND RAPIDS FORMATION: fine-grained quartzose sandstone; marine

CRAND RAPIDS FORMATION: fine-grained quartzose and feldspathic sandstone, laminated siltstone and silty shale; thin coaly beds; deltaic to marine

CLEARWATER FORMATION: dark grey fossiliferous silty shale, laminated siltstone, fine-grained cherty sandstone; glauconitic sandstone (WABIS-KAW MEMBER) near base; marine

LOON RIVER FORMATION: dark grey fossiliferous silty shale and laminated siltstone, ironstone nodules and partings; marine

BASAL CRETACEOUS, Wood Buffalo Park: calcareous quartz sandstone, equivalent in position to McMURRAY FORMATION of Athabasca River region

region

15a McMURRAY FORMATION: crossbedded quartzose sandstone and siltstone; oilimpregnated with grey silty shale interbeds in upper part; deltaic

DEVONIAN

GROSMONT FORMATION: grey fine-grained granular partly vuggy dolomite;
marine

MIKKWA FORMATION, upper member: brownish-grey to olive green, purplish-red mottled, fine-grained limestone, dolomitic limestone and shaly limestone; lower member: dark grey to brownish-grey fine-grained limestone and dolomitic limestone; marine

IRETON FORMATION: greenish-grey shale and calcareous shale and siltstone; marine

WATERWAYS FORMATION: grey and greenish-grey shale and argillaceous limestone units alternating with grey and greyish-brown fine-grained and clastic limestone units; marine

12a HAY RIVER FORMATION: greenish-grey calcareous shale, argillaceous silt-

CARIBOU MEMBER, SLAVE POINT FORMATION: grey and brown fine- to medium-grained limestone and dolomitic limestone; minor shale, gypsum; marine

10c FORT VERMILION MEMBER, SLAVE POINT FORMATION: white and grey gypsum, argillaceous and dolomitic gypsum; minor anhydrite; some fine-grained limestone, dolomitic limestone and green shale; evaporitic

10b MUSKEG FORMATION: white and grey gypsum, anhydrite; minor dolomite, salt and shale, evaporitic

NYARLING FORMATION: white and grey gypsum, argillaceous and dolomitic gypsum; some fine-grained limestone and dolomitic limestone; anhydrite, minor dolomite and salt, greenish shale and limestone in subsurface;

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KEC RIVER FORMATION, upper member: grey and brown fine- to medium-grained limestone, thin-bedded to massive, locally richly fossiliferous and coarse grained; grey medium-grained dolomite, locally vuggy; lower member: fine-grained brown dolomite and laminated dolomite, locally vuggy; minor limestone; unit locally much brecciated and recemented;

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CHINCHAGA FORMATION: white and grey gypsum, argillaceous and dolomitic gypsum and anhydrite; some dolomite, dolomitic limestone, red and green shale; some salt in subsurface; evaporitic

and green shale; some salt in subsurface; evaporitic

FITZGERALD FORMATION: grey and brown fine- to medium-grained dolomite to sandy dolomite and dolomitic limestone, locally vuggy; grading down into sandstone, arkose and breccia (LA LOCHE FORMATION);

MIDDLE DEVONIAN (undivided) includes METHY FORMATION: brown and buff massive porous dolomite, brown to grey thin-bedded dolomite, dolomitic limestone, minor anhydrite and gypsum; McLEAN RIVER FORMATION and leached PRAIRIE EVAPORITE FORMATION: gypsum, anhydrite, grey-green silty and dolomitic shale, minor dolomite; marine to evaporitic

PRECAMBRIAN

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ATHABASCA FORMATION: white, grey and red, medium- to coarse-grained sandstone, flat-bedded to strongly crossbedded; local conglomerate and shale beds

METASEDIMENTARY ROCKS: black and grey quartzite and impure quartzite, biotite schist, phyllite; minor sericite and hornblende schist, phyllonite, amphibolite; variable amounts of pegmatite and granite

amphibolite; variable amounts of pegmatite and granite

GRANITE GNEISS: predominantly biotite and hornblende granite gneiss;
minor amphibolite, schist, foliated granite, pegmatite and other gneisses

2b GRANITE: pink to grey, massive to foliated biotite, hornblende and leucocratic granite; local granite pegmatite

UNDIVIDED GRANITIC PLUTONIC ROCKS: including biotite granite, porphyroblastic and porphyritic granites; some granite gneiss and metasedimentary rocks

PORPHYROBLASTIC GRANITES: foliated to massive, white, grey and pink biotite granite and gneissic granite, quartz diorite, with feldspar porphyroblasts; variable amounts of metasedimentary rocks and granite gneiss

Ceological boundary
Outcrop location
Loose slab, probably from nearby unexposed bedrock
Spring
Thick drift
Sinkhole area
Surface contour (contour interval 500 feet)

Geology by R. Green and G. B. Mellon, 1962 M. A. Carrigy and R. Green, 1965

supplemented by published maps and reports of the Geological Survey of Canada and the Research Council of Alberta. Outcrop data for Dunvegan and Shaftesbury Formations in the vicinities of Cardinal and Bison Lakes supplied by O. Tokarsky, 1966, and for Precambrian areas south of Lake Athabasca by L. A. Bayrock and J. D. Codfrey, 1969.



North

Wapti formation

Puskwaskau Formation

Kaskapau Formation

Smaky
Group

Durvegan Formation

Durvegan Formation

Shaftesbury Formation

Shaftesbury Formation

Shaftesbury Formation

Loon River Formation

Loon River Formation

Devonion

Devonion

Devonion

North

Southeast

Southeast

Southeast

Southeast

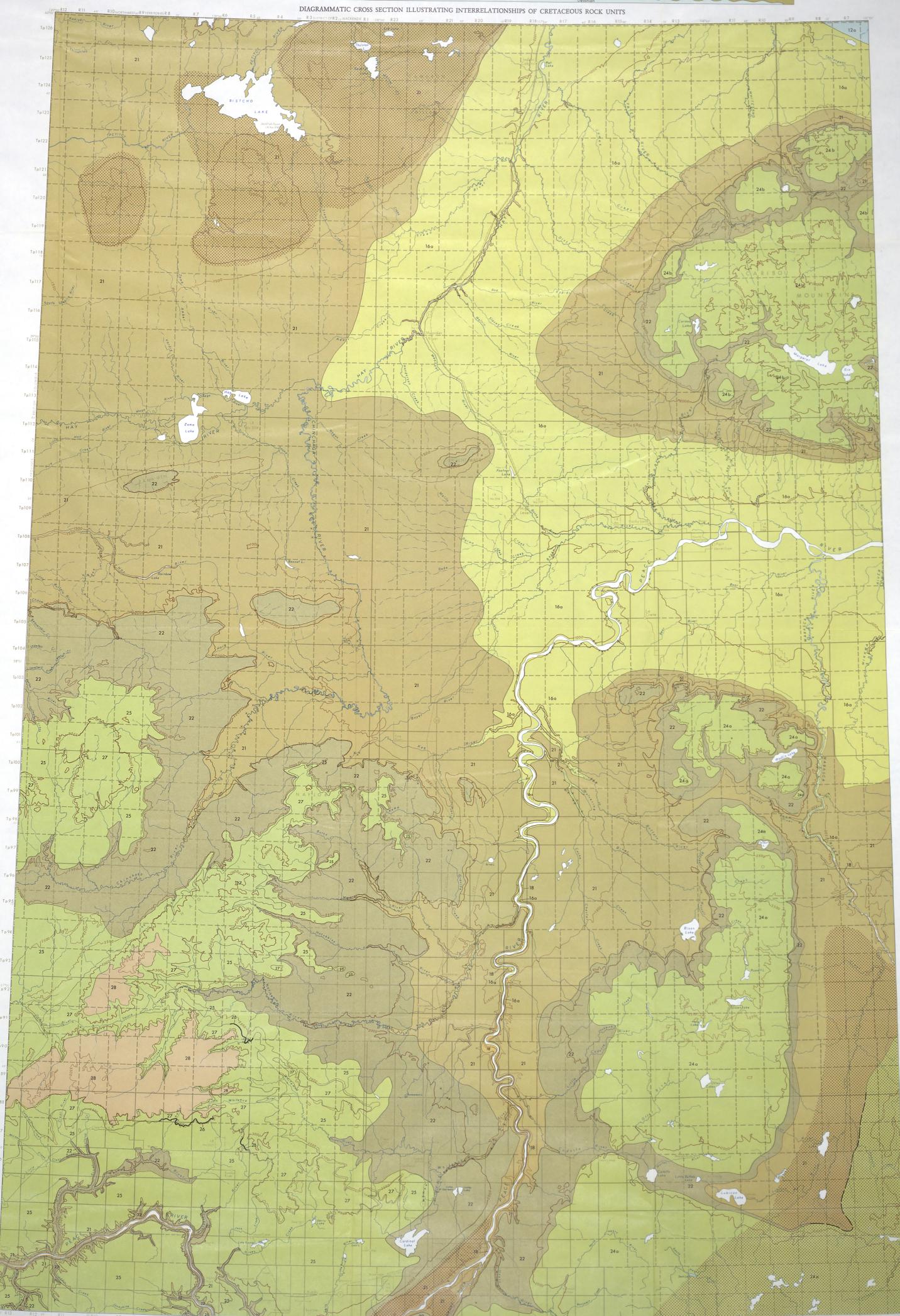
Southeast

Southeast

Southeast

Southeast

Southeast



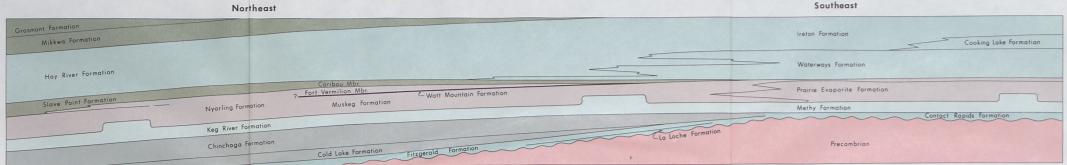
WEST HALF — FOR DESCRIPTIVE NOTES SEE EAST HALF

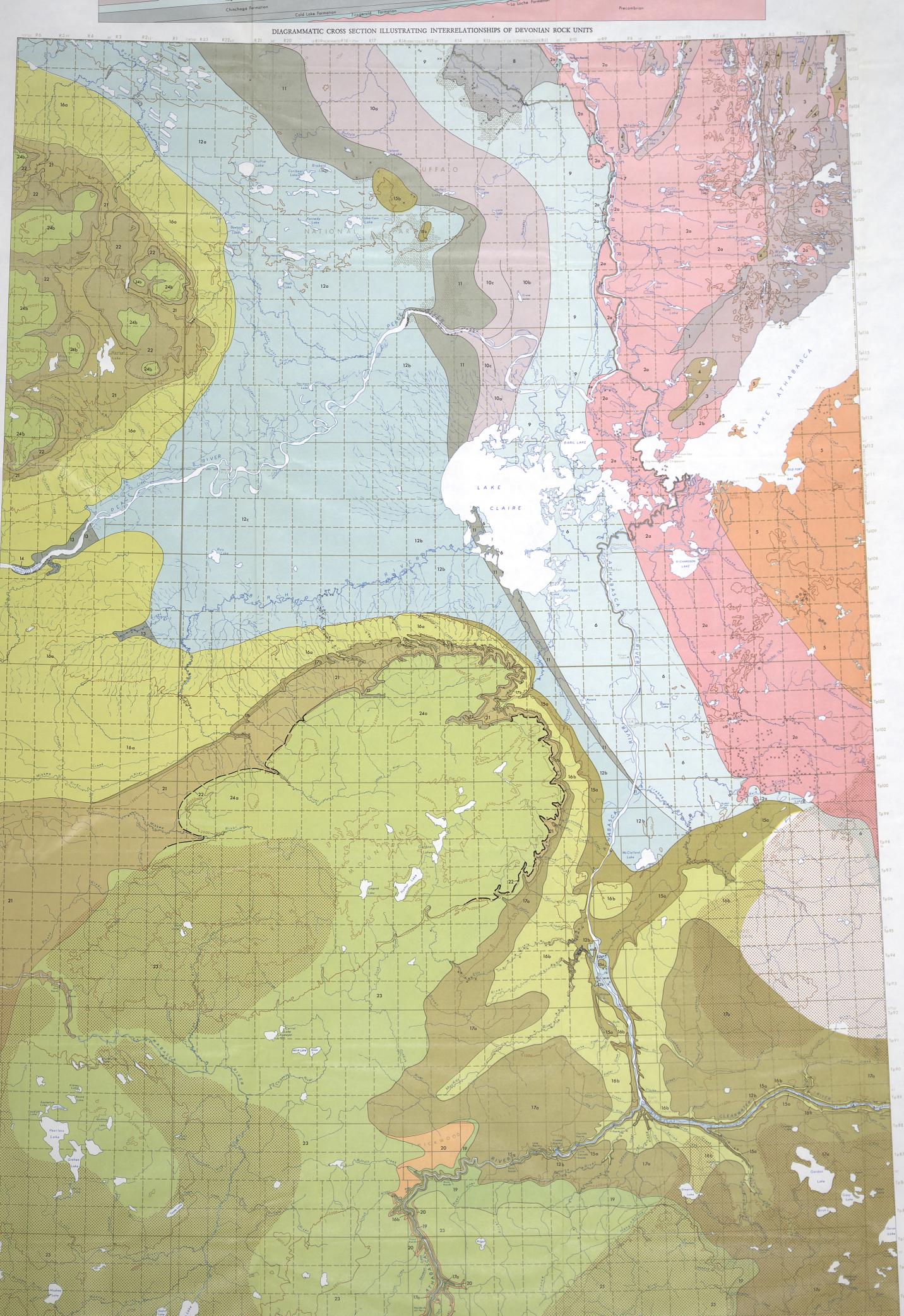
SCALE 1:500,000

1 Inch to 8 Miles Approximately
10 20 30 40

10 20 30 40

Published, 1970





DESCRIPTIVE NOTES

PHYSIOGRAPHY

PHYSIOGRAPHY

Alberta north of 56 degrees latitude covers an area of approximately 50,000 square miles, divisible into two major physiographic units: the Canadian Shield, which occupies about 6,000 square miles in the northeast corner of the province, and the remaining area, which forms part of the Interior Plains.

The Canadian Shield, composed of igneous, metamorphic and sedimentary rocks of Precambrian age, is an area of low but irregular relief with numerous lakes and, in the northern part, a high proportion of bedrock outcrop. South of Lake Athabasca bedrock outcrops are scarce, and the area inferred to be underlain by Precambrian rocks is largely covered by unconsolidated deposits of Pleistocene and Recent ages.

The Plains, a region of low to moderate relief, is underlain by gently dipping rocks of Devonian and Cretaceous ages. A dissected plateau, the area consists of a series of generally steep-sloped, concordant uplands (Clear, Chinchaga, Cameron, Buffalo Head and Thickwood Hills; Caribou and Birch Mountains) which rise one to two thousand feet above the general level of the gently undulating lowlands adjacent to the Hay, Peace and Athabasca Rivers. Locally, the topography is modified by the effects of Pleistocene glaciation, and the entire region is covered by a veneer of unconsolidated glacial, aeolian and Recent deltaic deposits. Most of the region is forested or covered by poorly drained muskeg and swamp deposits, except for partly cultivated areas in the lowlands adjacent to the Peace River and the McKenzie Highway. Due to the thick drift deposits and vegetative cover, bedrock exposures are confined mainly to the valleys of the larger rivers and tributary streams, and along the slumped margins of the upland areas.

PRECAMBRIAN

The Precambrian Shield area comprises a complex of igneous and metamorphic rocks ranging in age from 1.7 to 2.3 billion years and cut by a number of major and minor fault structures. North of Lake Athabasca an average of 60 per cent of the area is rock outcrop. South of the lake much less outcrop is found: the most extensive outcrop area, between Reid Creek and Marguerite River, has bedrock on 25 per cent of the surface.

The GRANITE GNEISS complex contains the oldest rocks—a heterogeneous mixture of biotite and hornblende gneisses with minor bands of metasedimentary rocks and amphibolites. The METASEDIMENTARY ROCKS represent either remnants of paragneissic parent material or enclosed bodies of younger sedimentary material. A metasedimentary band may contain pure to impure quartzite, amphibolite and biotite, hornblende and sericite schists.

and sericite schists.

PORPHIROBLASTIC GRANITES, both gneissic and biotitic, form a distinctive rock group north of Lake Athabasca. These granites have developed from and are intimately associated with metasedimentary and metavolcanic rocks. A second group of GRANITES—massive to foliated biotite, hornblende and leucocratic rocks —intruding other rock types was formed subsequent to the main tectonic deformation of the region. Cataclastic deformation has resulted in the formation of mylonite zones over a mile wide and tens of miles long.

The ATHABASCA FORMATION, predominantly an orthoquartzitic sandstone, is poorly exposed in the area south of Lake Athabasca, adjacent to the Saskatchewan border. Both well-cemented and poorly cemented phases exist, the latter being extremely difficult to differentiate from unconsolidated sands glacially derived from Athabasca Formation parent material.

DEVONIAN

The Middle Devonian FITZGERALD FORMATION, outcropping at scattered points along Slave River, consists of dolomitic limestones and dolomites grading down into the sandstones and regolithic igneous detritus of the La Loche Formation which rests on Precambrian rocks. Up to 19 feet of La Loche and 26 feet of Fitzgerald

or the La Loche Formation which rests on Precambrian rocks. Up to 19 feet of La Loche and 26 feet of Fitzgerald Formation are exposed.

The CHINCHAGA FORMATION, exposed extensively in the escarpment and sinkhole area southwest of Salt River and intermittently along Salt and Slave Rivers, consists of gypsum and minor anhydrite, shale and dolomite up to 100 feet being exposed. In the subsurface westward the full thickness of 250 feet is attained; however, the unit is largely removed by solution east of the escarpment the Slave River outcrops consisting of the basal shales and minor gypsum. Saline springs, near the Salt River, deposit salt derived from solution of the underlying, unexposed Cold Lake Formation.

The KEG RIVER FORMATION, forming the caprock along the escarpment west of Salt River, consists of limestones and dolomites up to 200 feet thick. The outcrop area also extends along Salt and Slave Rivers, where solution of underlying evaporite drops the unit by 200 feet. Extensively brecciated beds are exposed in places along Salt and Slave Rivers.

The MIDDLE DEVONIAN (undivided) map-unit covers: (1) a large area south of Lakes Claire and Athabasca believed to be underlain by Middle Devonian strata, and (2) the Paleozoic outcrop belt in the Clearwater River valley. Outcrops in area (1) are confined to Firebag and Marguerite Rivers and environs of fessilferous dolomite slabs at several other locations suggest the proximity of bedrock. All exposures appear to be of Methy Formation (Keg River-equivalent) strata—reefal, algal and fine-grained dolomites, locally fossilferous and spore bearing, and with local calcareous brecas. In the Clearwater River valley, where the Methy unit reaches a thickness of 200 feet, up to 80 feet of dolomite outcrop at the various falls. The overlying McLean River Formation shales and silstones and remnants of the leached Prairie Ecaporite (gypsum, anhydrite and evaporitic dolomite and shale) should be represented in the outcrop area, but only salt springs indicate the presence of ev

Ecaporite (gypsum, anhydrite and evaporitic dolomite and shale) should be represented in the outcrop area, but only salt springs indicate the presence of evaporites in the subsurface.

The NYARLING FORMATION is the map-unit containing equivalents of the subsurface Muskeg (evaporite) and Watt Mountain (shale-siltstone) Formations and the Fort Vermilion (evaporite) Member of the Slave Point Formation, Muskeg Formation, the northern equivalent of the Prairie Evaporite, is 500 to 750 feet thick west of the fifth meridian, and consists of anhydrite and dolomite with up to 250 feet of halite. Watt Mountain Formation consists of 20 to 50 feet of greenish shale and siltstone, limestone and breccia. No outcrops of either rock unit were found; the rock unit boundaries are therefore somewhat hypothetical. FORT VERMILION MEMBER consists of up to 120 feet of anhydrite in subsurface sections; up to 75 feet of gypsum probably of this unit outcrop along the Peace River at Peace Point. The unit thins northward (to 40 feet in LO.E. Jack Lakes 8-17-120-1 W.5 well) and disappears southward away from the lower Peace River region. Solution of the evaporite succession, indicated by extensive sinkhole areas, has reduced Nyarling strata to an estimated 200-foot maximum along the subcrop zone in Wood Buffalo Park.

The SLAVE POINT FORMATION (CARIBOU MEMBER), consisting of brownish dolomitic limestone and limestone, is well exposed (70 feet) at Peace Point. Small outcrops in sinkholes in Wood Buffalo Park and on Lake Claire constitute the only other exposures. The unit thins from over 120 feet at 60°N to 70 feet at Peace Point (by facies change), southward to less than 10 feet in the Birch Mountains (possibly by onlap).

onlap).

The HAY RIVER FORMATION, approximately 1,300 feet thick where complete, is a succession of greenish-grey calcareous shale with fine-grained limy stringers. Limestone beds become more common south of the Peace River as the lower 500 to 600 feet of the unit change facies to the Waterways-Beaverhill Lake Formations; the facies boundary is arbitrarily placed along the lower Peace River. The name Ireton Formation is used for the Hay River-equivalent shales between the Waterways and the overlying Mikkwa Formations. The WATERWAYS FORMATION, a succession of greenish-grey calcareous shale and argillaceous limestone alternating with grey and greyish-brown, fine-grained clastic limestone is 600 to 700 feet thick. It is best exposed along the Athabasca River north of Fort McMurray.

The MIKKWA FORMATION, a 400-foot thick succession of limestone and dolomitic limestone with minor shale, grey, green and reddish in color, outcrops along the Peace River in the Vermilion Chutes area and on Harper Creek to the south. The overlying GROSMONT FORMATION is a reefal, coarsely vuggy, petroliferous dolomite 400 feet thick, of which 50 feet outcrop on the Peace River in the Vermilion Chutes area.

CRETACEOUS

The McMURRAY FORMATION, lying unconformably on Devonion carbonate rocks, is well exposed along the valleys of the Athabasca River and its tributary streams. The unit is from 150 to 300 feet thick, thinning to the north and west beneath the Birch Mountains. It is composed mainly of deltaic sediments in the form of thick crossbedded, oil-impregnated, quartz sands with interbeds of laminated pale grey silt and silty shale. Locally, thin tongues of marine shale are present in the upper part, which is conformable with the overlying Clearwater Formation. Beds mapped as BASAL CRETACEOUS are those of inferred Cretaceous age equivalent in stratigraphic position to the McMurray Formation poorly exposed in the lowlands east of the Caribou Mountain Wood Buffalo Park. They consist of scattered outcrops of hard, calcite-cemented, quartz sandstone overly Devonian carbonate rocks. Their areal extent is conjectural owing to lack of exposures.

The LOON RIVER FORMATION underlies a large region in the lowlands adjacent to the Peace and Hay Rivers in the northern part of the map-area. It is from 600 to 800 feet thick, extending to the east around the north flank of the Birch Mountains, where it is contiguous with the Clearwater Formation. The unit consists of marine dark grey, fossiliferous shale and laminated siltstone with scattered nodules and thin beds of sideritic ironstone. The upper boundary is gradational with the Peace River Formation in the south and conformable with the Shaftesbury Formation in the north. The CLEARWATER FORMATION, consisting of marine dark grey fossiliferous shale, laminated siltstone and thin beds of fine-grained cherty sandstone, is exposed along the Athabasca River and its tributaries. It is about 350 feet thick in the Fort McMurray area, thickening at the expense of the overlying Grand Rapids Formation towards the north end of the Birch Mountains. The base of the formation is marked by a thin glauconitic sandstone called the Wabiskaw Member, and the upper boundary is gradational with the Crand Rapids Formation.

The PEACE RIVER FORMATION is exposed along the Peace and Wabasca Rivers south of township gradational with the Grand Rapids Formation.

The PEACE RIVER FORMATION is exposed along the Peace and Wabasca Rivers south of township gost it is divisible into three members—Northevoin (lower) Member: 50 to 75 feet of thin-bedded, fine-grained, glauconitic, feldspathic sandstone and silstone, grading at the base into silty shale; Harmon (middle) Member: marine dark grey shale; Cadotte (upper) Member: fine-grained, quartzose, cherty sandstone with thin coal and silty shale interbeds in the upper part. To the north the unit passess laterally into marine shale in the upper part of the Loon River Formation; to the east in the subsurface, it is contiguous with the Grand Rapids Formation. The GRAND RAPIDS FORMATION underlies much of the thick drift-covered area adjacent to the Saskatchewan border, extending west to outcrop along the valley of the Athabasca River and the lower slopes of the Birch Mountains. South of Fort McMurray, the unit is about 350 feet thick, consisting of unconsolidated fine-grained quartzose and feldspathic sandstone and siltstone interbedded with marine silty shale. The lower 100 to 150 feet is composed of glauconitic cherty sandstone conformable worthying the silty shales of the Carawater Formation. The formation thins to the north along the margin of the Birch Mountains, where it is contiguous with the ALICE CREEK TONGUE. This name is proposed for a homogeneous, fine-grained quartzose sandstone 50 feet or more in thickness exposed along the north slope of the Birch Mountains. In sharp but conformable contact with the overlying Shaftesbury Formation, the unit is considered equivalent to the upper beds of the Grand Rapids Formation.

The PELICAN and JOLI FOU FORMATIONS are exposed along the valley of the Athabasca River

Grand Rapids Formation.

The PELICAN and JOLI FOU FORMATIONS are exposed along the valley of the Athabasca River southwest of Fort McMurray. The lower unit, the Joli Fou, is up to 100 feet thick, consisting of marine dark grey, fossiliferous shale with glauconitic silt interbeds in the upper part. The Pelican, 40 to 50 feet thick, consists of fine-grained quartzoes sand with interbeds of glauconitic silty shale in the lower part. The unit appears to shale out northwest of Fort McMurray; equivalent beds in the Birch Mountains, if present, are included in the based next of the Schechour Formation. basal part of the Shaftesbury Formation.

basal part of the Shaftesbury Formation.

The SHAFTESBURY FORMATION underlies much of the lowland area adjacent to the Peace, Hay and Chinchaga Rivers in the western part of the map-area, extending eastward around the lower slopes of the Caribou Mountains, Buffalo Head Hills and northern margin of the Birch Mountains. The unit is from 800 to 1,000 feet thick, consisting of marine dark grey to black, highly fissile shale, scattered thin bentonite beds and abundant concretionary ironstone. Thin, laminated, fish scale-bearing siltstones are present in the lower part, and a fine-grained cherty sandstone interval, 20 to 30 feet thick about 250 feet above the base of the formation, is exposed along the south flank of the Caribou Mountains. The Shaftesbury is prone to large-scale slumping in the form of landslides and mudflows, which are associated in places with spontaneous combustion of shales rich in organic matter. The upper boundary with the Dunvegan Formation appears conformable and gradational from subsurface data; in the southeast part of the map-area, where the Dunvegan is absent, equivalent beds are included in the Labiche Formation. in the Labiche Formation.

in the Labiche Formation.

The DUNVEGAN FORMATION underlies the lower slopes of the uplands in the southwest part of the map-area, rising to the north and east to form a narrow outcrop band about the upper slopes of the Caribou Mountains and the northern margin of the Birch Mountains. Exposures are generally poor and discontinuous, consisting mainly of friable, pale grey, fine-grained, feldspathic sandstone with scattered hard calcareous beds, laminated carbonaceous slitstone and dark grey silty shale. Although fossil remains are scarce, the beds appear to be of deltaic origin. The unit thins from 500 feet in the southwest to less than 200 feet in the Caribou and Birch Mountains, grading laterally into marine shale of the Labiche Formation in the southeast part of the map-area.

Birch Mountains, grading laterally into marine shale of the Labiche Formation in the southeast part of the map-area.

The term SMOKY GROUP is used for a succession of marine, predominantly shaly beds overlying the Dunvegan Formation in the southwest part of the map-area. Where the Dunvegan Formation is absent by facies change in the southeast, the rocks are included with the lithologically similar Shaftesbury Formation as part of the LaBICHE FORMATION. West of the Peace River, the Smoky Group is divisible into three units. The lower unit, the KASKAPAU FORMATION, is about 500 feet thick near the southwest corner of the map-area, thinning to less than 200 feet to the northeast. The lower 50 to 150 feet of strata comprise a succession of fine-grained quartzose sandstone, dark grey silty shale and thin beds of ferruginous solitic ironstone conformably overlying the Dunvegan Formation. The upper 150 to 400 feet consist of marine dark grey, silty shale gradational at the top with the overlying Bad Heart Formation. The BAD HEART FORMATION, from 0 to 30 feet thick, is composed of dark brown to green, ferruginous, solitic sandstone and mudstone with thin layers of siderite pebble conglomerate. Drilling data show that the unit consists of northwest-trending linear bodies of rock that form substantial reserves of low grade iron ore in the Clear Hills area. The PUSKWASKAU FORMATION underlies the upper slopes of the highlands in the southwest part of the map-area. It is poorly exposed, consisting of marine dark grey, fossiliferous shale 300 to 400 feet thick in the Clear Hills. Seattered exposures of marine dark grey, shale are present near the headwaters of small streams flowing across the upper surface of the Caribou Mountains and northern part of the Buffalo Head Hills. Overlying the Dunvegan Formation and attaining an estimated maximum thickness of at least 400 feet, the shale is correlated with the Smoky Group to the south on the basis of lithology and stratigraphic position.

The WAPITI FORMATION caps the Clear and pa

SURFICIAL DEPOSITS

Most of the map-area is covered by unconsolidated glacial, fluvial and lacustrine deposits ranging in thickness from a few to several hundred feet. The oldest deposits consist of quartzitic and cherty gravels of fluvial origin overlying bedrock in the Clear Hills. Similar deposits are probably present capping other plateau remnants in the region. Quartzitic gravels containing minor amounts of granite and volcanic rock pebbles also underlie glacial drift of the Plains area south and east of the Clear Hills, being found at several elevations. These deposits are post-Cretaceous and preglacial, although at least some of the fluvial sand and gravel successions preserved in preglacial river valleys are Pleistocene in age.

Glacial sediments of Pleistocene age cover most of the Plains and much of the Canadian Shield. Of variable thickness and lithology, the deposits consist mainly of clayey to sandy till and proglacial lacustrine silts and clays. Thick deposits of outwash sand and gravel also are common, especially in the eastern part of the map-area, south of Lake Athabasca.

Unconsolidated fluvial, lacustrine, and aeolian sediments of Recent age are scattered throughout the map-area. The thickest accumulation is at the west end of Lake Athabasca, in the delta formed by the Peace and Athabasca Rivers. Large sand dunes are present in the flat terrain east of the Caribou Mountains and also south of Lake Athabasca; smaller dune fields are present elsewhere in the map-area.

OUTCROPS

OUTCROPS

All known outcrops are indicated by symbol except in the following areas:

(1) the Peace River valley between township 82, range 13, and township 81, range 6, west of 6th meridian;

(2) the Peace River valley between township 81, range 24, and township 94, range 20, west of 5th meridian;

(3) the Athabasca River valley between township 81, range 17, and township 97, range 11, west of 4th meridian;

(4) the Clearwater and lower Christina River valleys;

(5) north of Lake Athabasca and east of the Slave River; and

(6) between Marguerite River and Reid Creek in townships 100-101, ranges 4-5, west of 4th meridian.

In these specific areas the extensive bedrock outcrops cannot be indicated without obscuring the rock unit boundaries.

GEOLOGICAL BOUNDARIES

Relatively precise geological boundaries are shown for the Precambrian Shield region north of Lake Athabasca, where approximately 60 per cent of the area is bedrock outcrop. Elsewhere outcrops are limited except along the larger river valleys; therefore, most rock unit boundaries shown are approximate or postulated, and are based on outcrop data supplemented by exploratory well information. Rock-unit boundaries are projected to the land surface except where subsurface data define subcrop limits, i.e. (1) McMurray Formation boundaries in the area adjacent to Fort McMurray and northward to McLelland Lake; (2) Clearwater-Gand Rapids boundary near Gregoire Lake, along the MacKay River and on the flanks of Muskey Mountain; (3) Joli Fou Formation boundaries in the Stony Mountain area; (4) Shaftesbury-Dunvegan boundary along the Peace River southeast of Grimshaw.

Where several rock units outcrop along steep-sided valleys of the Athabasca, Clearwater, Peace, and Wabasca Rivers and major tributary streams, unit boundaries are shown diagrammatically instead of accurately in order to illustrate the presence of all rock units; stream crossing points remain accurately located.

No rock units are shown in the upper Firebag River area as no control data exist.

DRIFT

The areas of thick drift, shown by stipple, are indicated by relatively scattered subsurface data, or are postulated from surface morphology (as in the upper Firebag River area). Drift, including sand and till units, up to 600 feet thick covers the north flank of Muskeg Mountain and much of the Stony Mountain area. Similar thicknesses indicate buried valleys beneath the present Lubicon, Loon, Muskwa and upper Wabasca Rivers, and adjacent to the Peace River southeast of Grimshaw. Much of the Cameron Hills area has substantial drift deposits that attain thicknesses of over 1,000 feet; the southwest portion of the Birch Mountains, the Buffalo Head Hills and the Caribou Mountains have thick drift accumulations locally.